

# Lecture 18: Volcano Deformation I



GEOS 655 Tectonic Geodesy

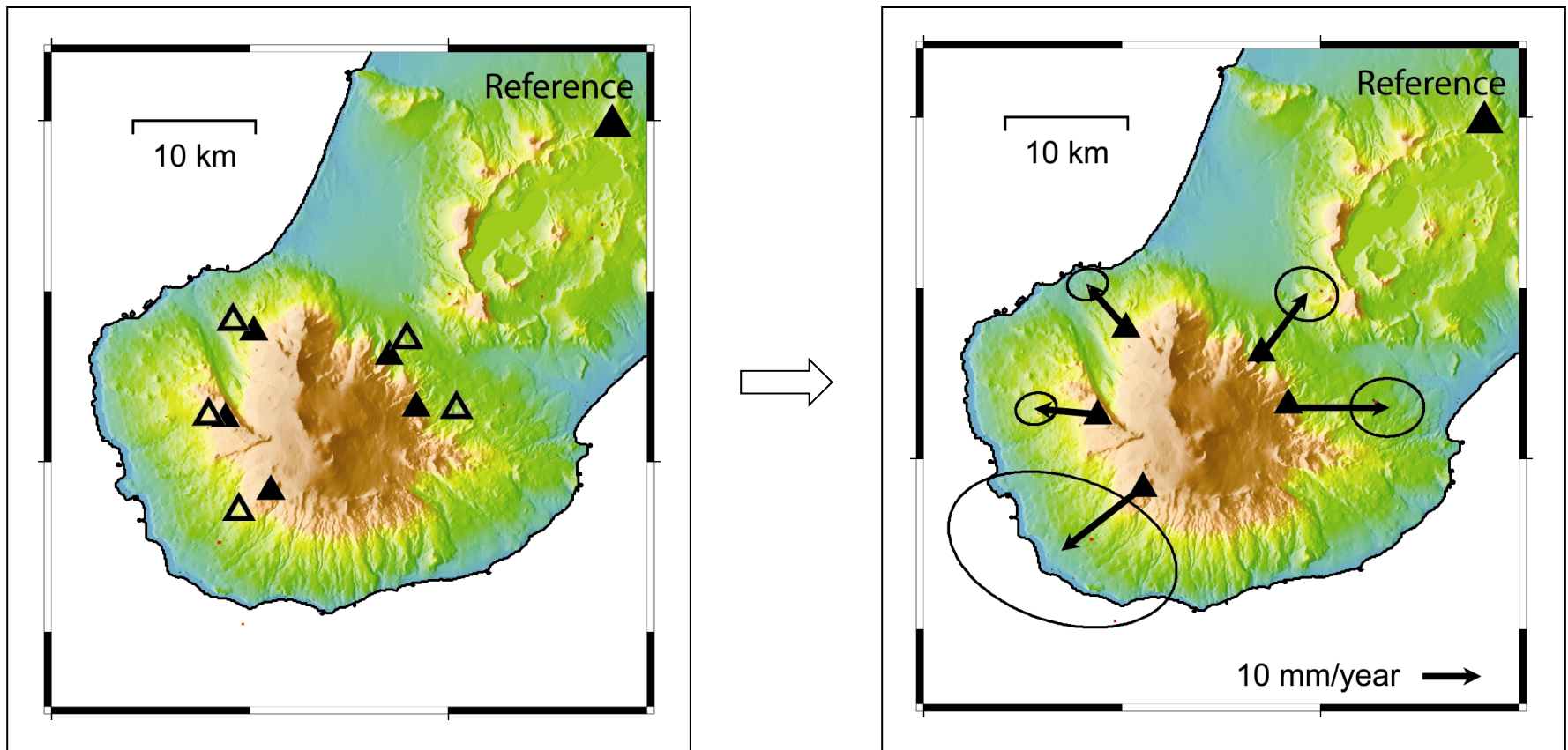
Jeff Freymueller

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# Cause of Volcano Deformation

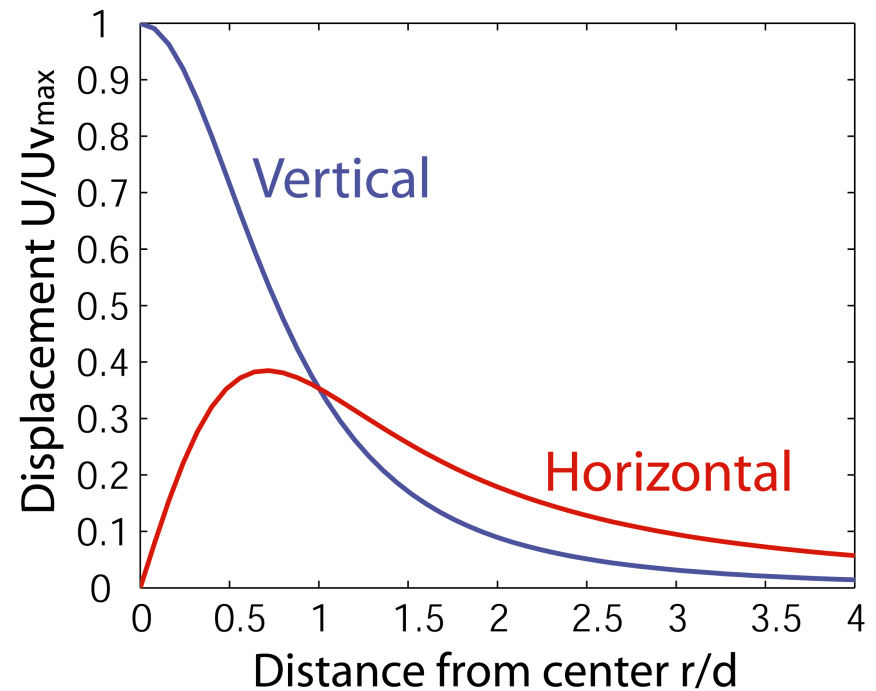
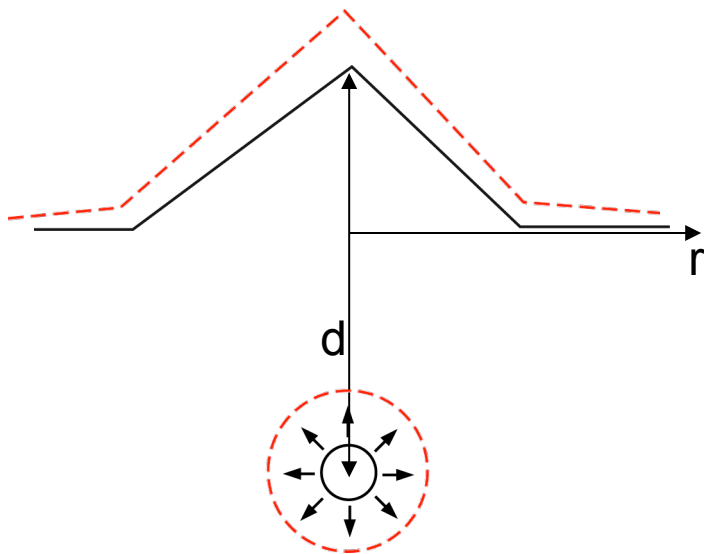
- Volcano deformation results from pressure changes caused by fluids.
  - Intrusions of magma in subsurface
  - Removal of magma to the surface during eruption
  - Pressure changes within hydrothermal systems
- Stresses and stressing rates from these sources can be very large.

# Example of Volcano Deformation (Westdahl, Aleutian Arc)



# Volcanic source models

Pressure point source  
(Mogi, 1958)

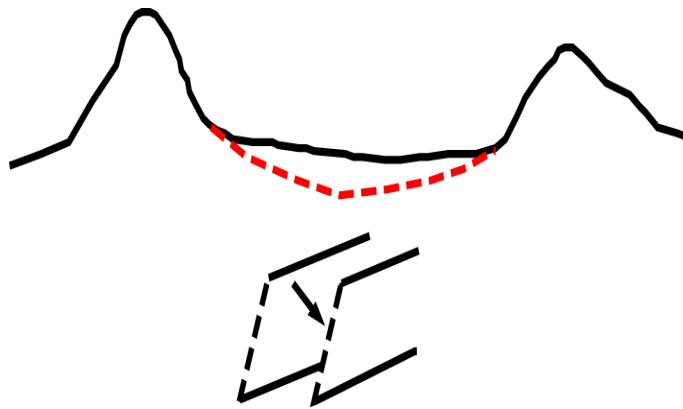


$U_{\text{max}}$ : max. vertical displacement

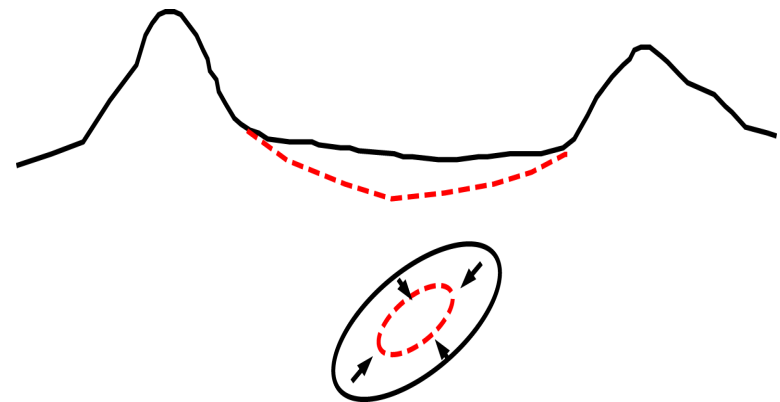
$d$ : source depth



Rectangular dislocation  
source (Okada, 1985)



Ellipsoidal pressure  
source (Yang, 1988)



Dike or sill, expand or contract    Pressure release, degassing

Model parameters:

Location, depth, strength, length, width, dip, strike

# Mogi model

Spherical pressure point source: **C**

$$U_z = \frac{(1-\nu)pa^3}{\mu} \frac{d}{(r^2 + d^2)^{3/2}}$$

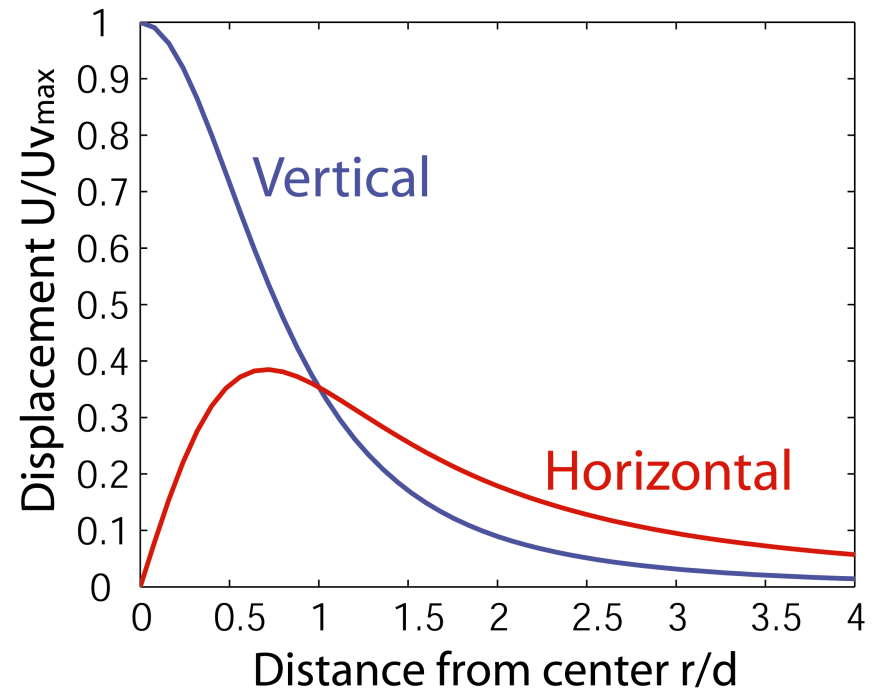
$$U_r = \frac{(1-\nu)pa^3}{\mu} \frac{r}{(r^2 + d^2)^{3/2}}$$

$$\Delta V = \pi pa^3 / \mu$$

p = pressurization

a = radius of sphere

$\Delta V$  = volume change



$U_{vmax}$ : max. vertical displacement

d: source depth

# Notes on Mogi Model

- Radial and vertical displacements are like the sine and cosine of the 3D displacement vector
  - Because displacements at surface are radially outward from pressure source
  - This is not true at all depths
- Magnitude of (3D) displacement vector decreases as  $1/\text{distance}$  from source
- The  $\Delta V$  is volume change in the chamber.
  - Cannot exactly equate to change in mass because of compressibility of magma.

# What is the $\Delta V$ ?

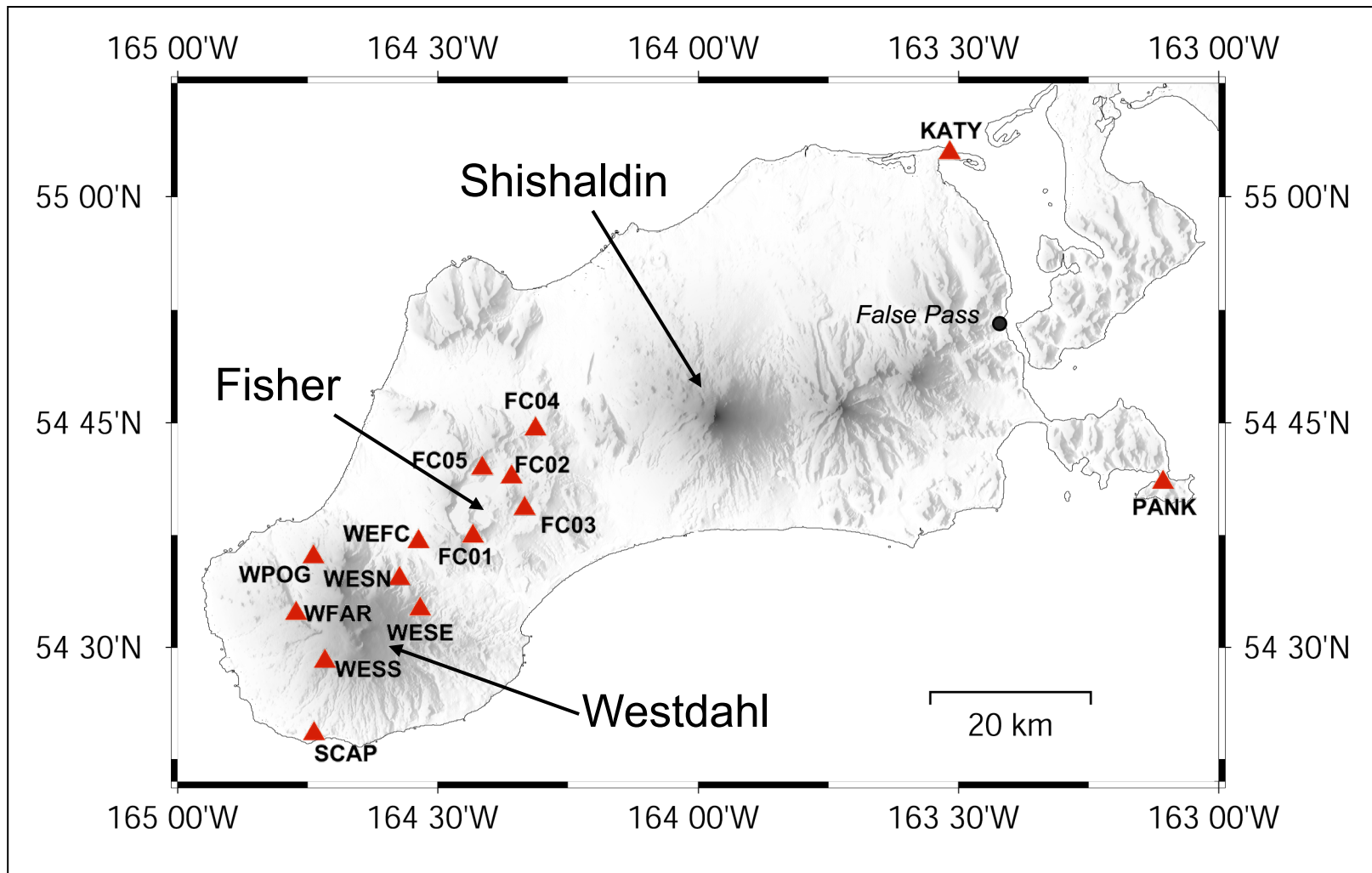
- The  $\Delta V$  in the previous slide is actually the volume change of the cavity, which is equivalent to a pressure change of the cavity.
- It may or may not reflect the volume of magma intruded
  - Incompressible magma: volume change of cavity equals volume change of magma
  - Compressible magma: what do you really mean by “volume” anyway? Volume is a function of pressure and mass
- Key points are that magma “volume” is model-dependent and also may or may not relate easily to mass of new magma

# Derivation of Mogi Model

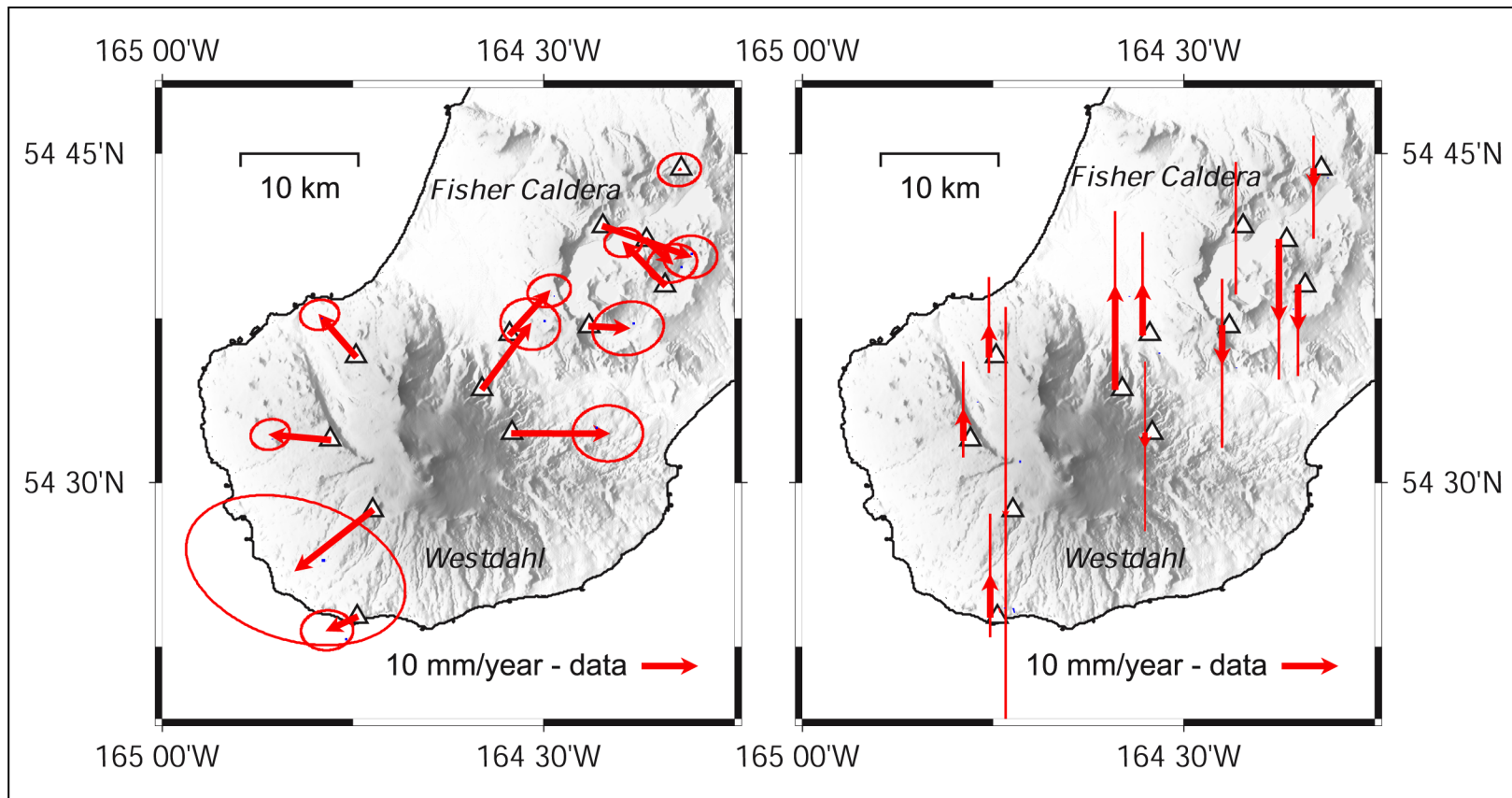
- Assume a spherical cavity of radius  $a$
- Assume stress on walls of cavity is isotropic, corresponds to pressure  $p$ .
- Solve differential equation in terms of displacement
  - Impact of free surface means that an iterative approach to finding the solution is necessary (outlined in Segall, Chapter 7).
- Point source approximation means  $a \ll d$ 
  - In practice, approximation is good for  $a < 0.5d$



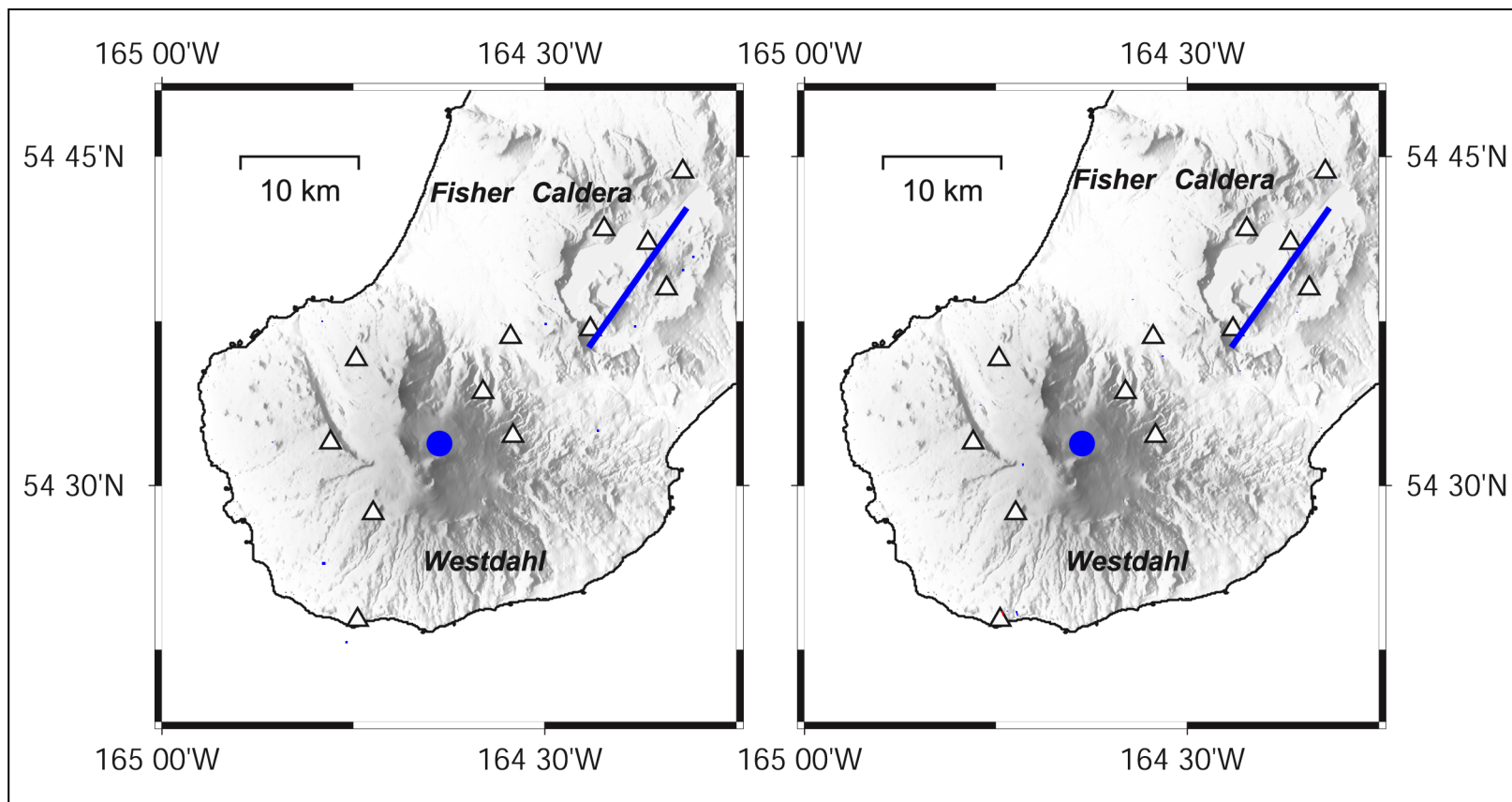
# GPS network



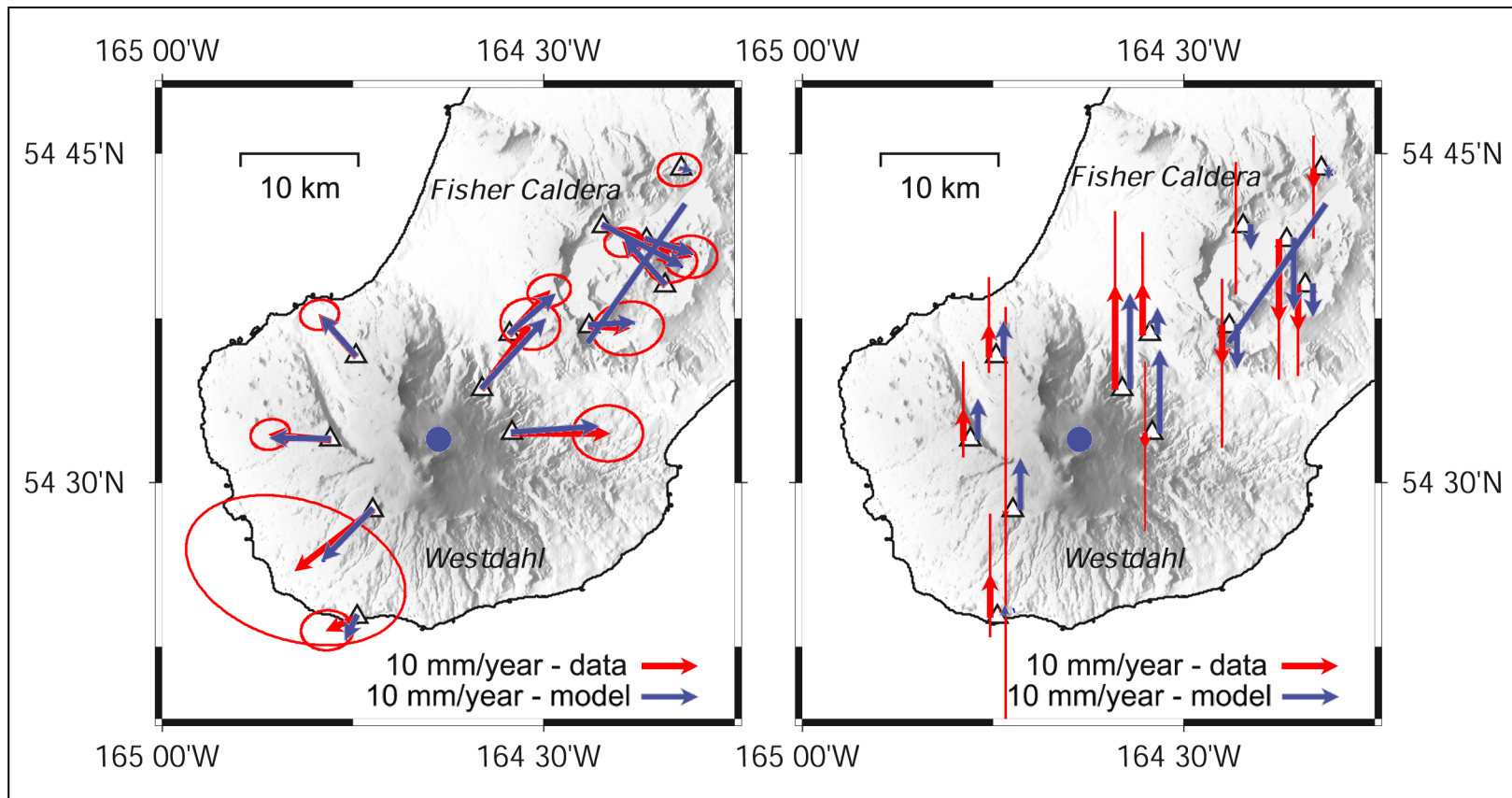
# GPS data (1998-2001)



# Model

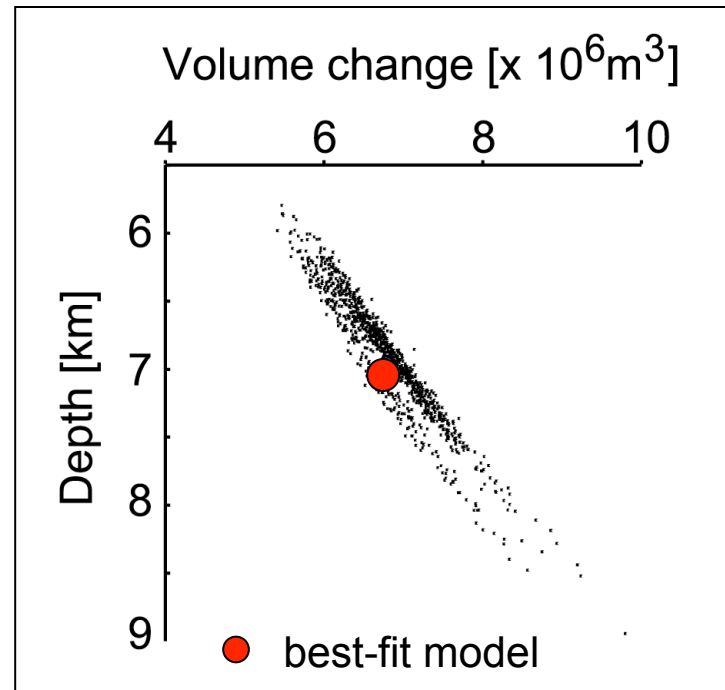
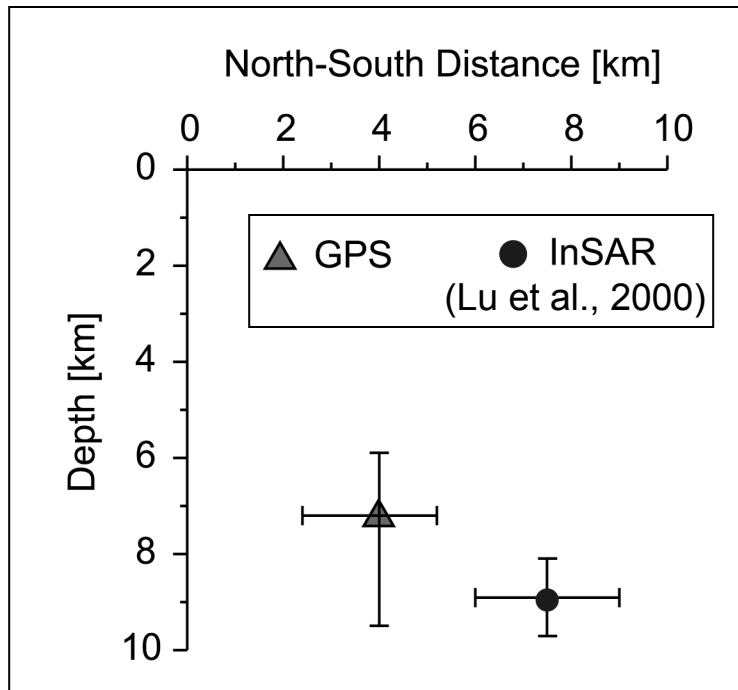


# GPS data and model



# Modeling results - Westdahl

- Point source east of Westdahl Peak
- Depth: 7.2 km (6 km below sea level)
- Volume change:  $6.7 \times 10^6 \text{ m}^3/\text{year}$



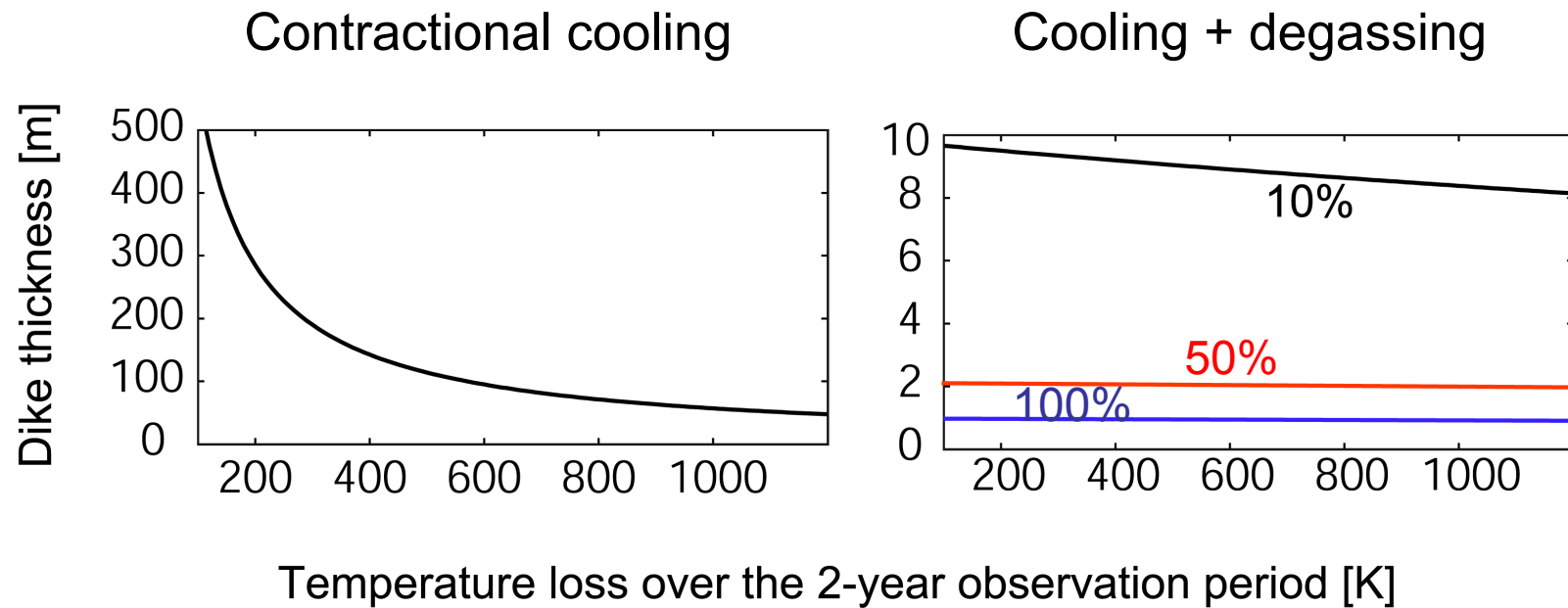


# Modeling results - Fisher

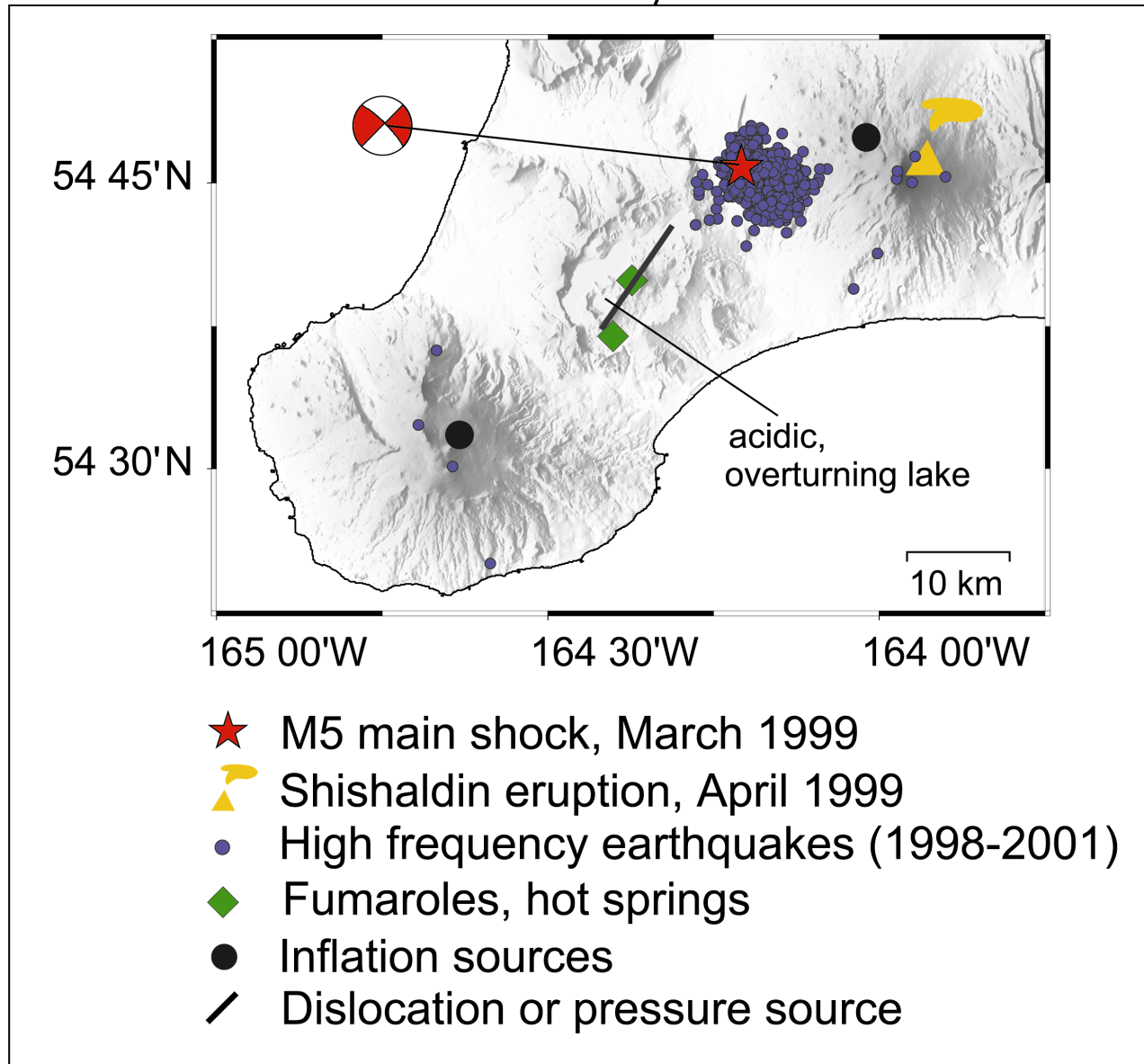
- Dike-like source along the long caldera axis

| Depth<br>(km) | Length<br>(km) | Width<br>(km) | Dip<br>(°) | Strike<br>(N°E) | Volume change<br>(x 10 <sup>6</sup> m <sup>3</sup> /year) |
|---------------|----------------|---------------|------------|-----------------|---|
| 1.5           | 13             | 0.5           | 80         | 35              | - 2.1   |

# Dike cooling model



# Volcanic, hydrothermal, and tectonic activity



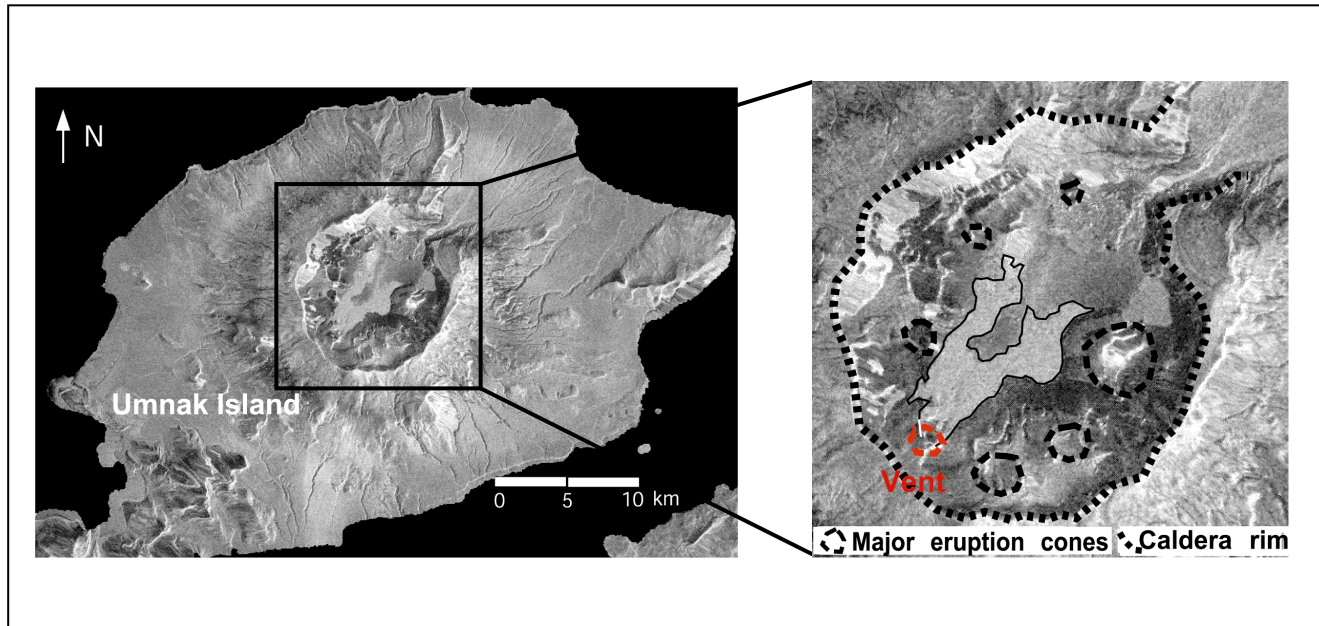
# Rapid and Time-Variable Inflation at Okmok Volcano, Aleutian Arc

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Yosuke Miyagi, Fumiaki Kimata

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Nagoya University*



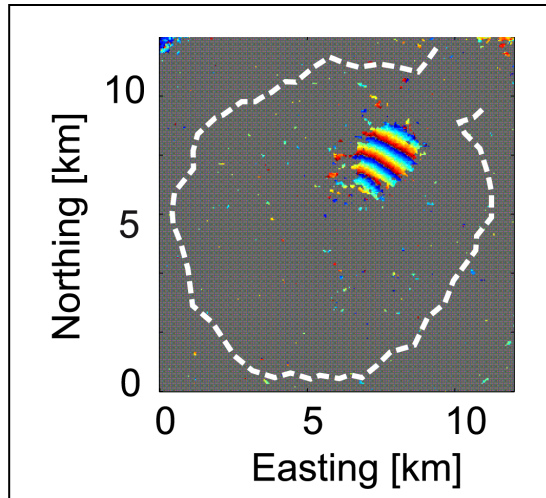
# Location



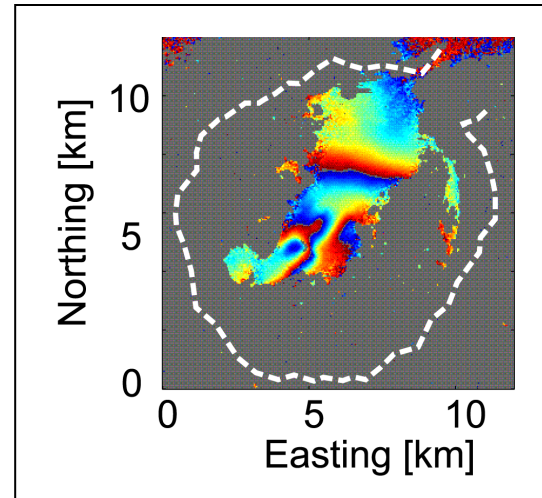


# Deformation interferograms

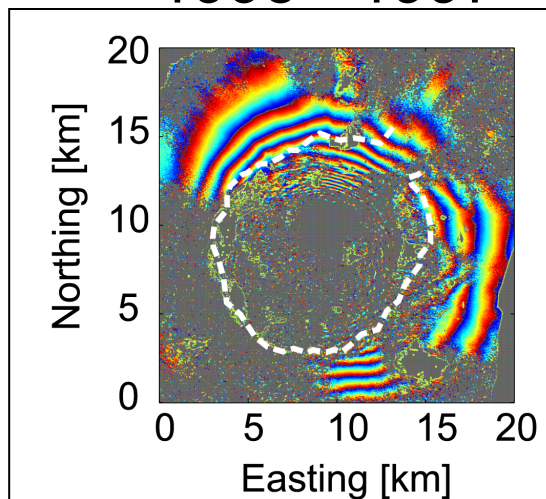
1992 - 1993



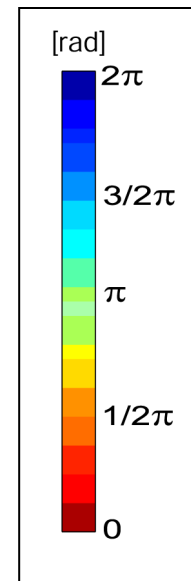
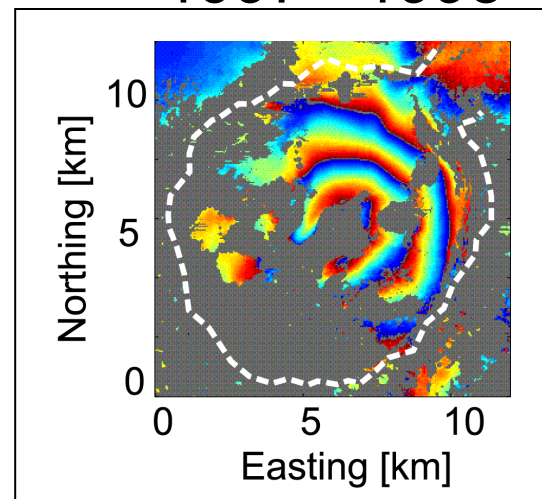
1993 - 1995



1995 - 1997

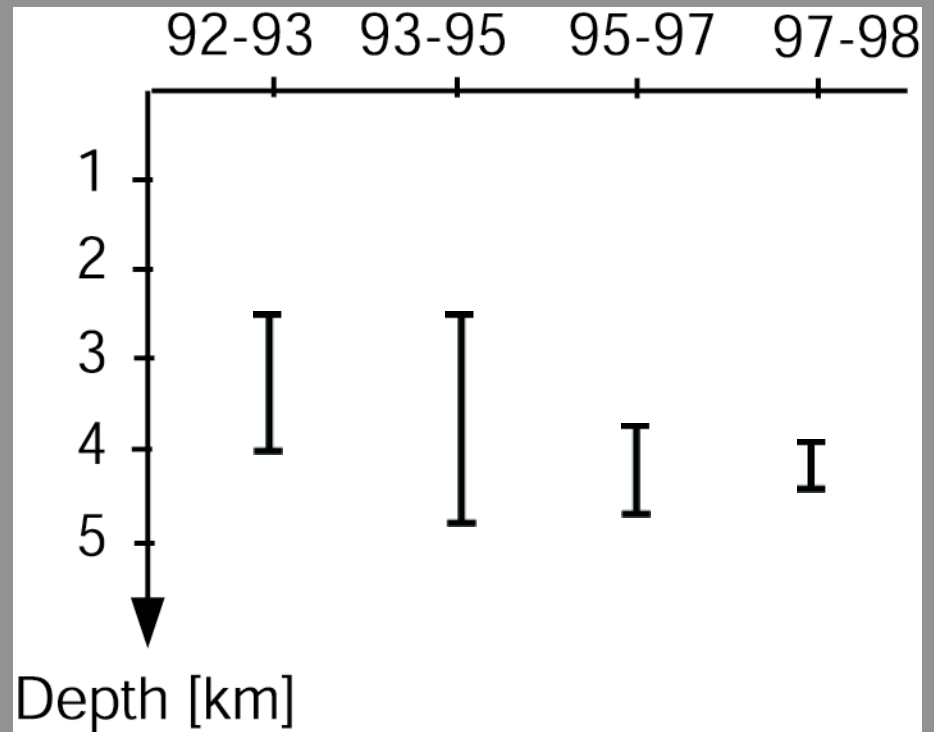
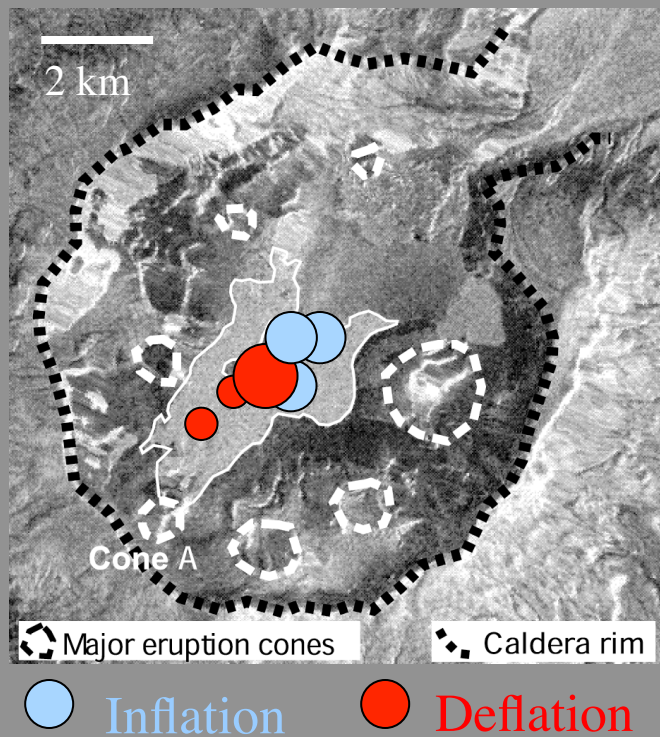


1997 - 1998



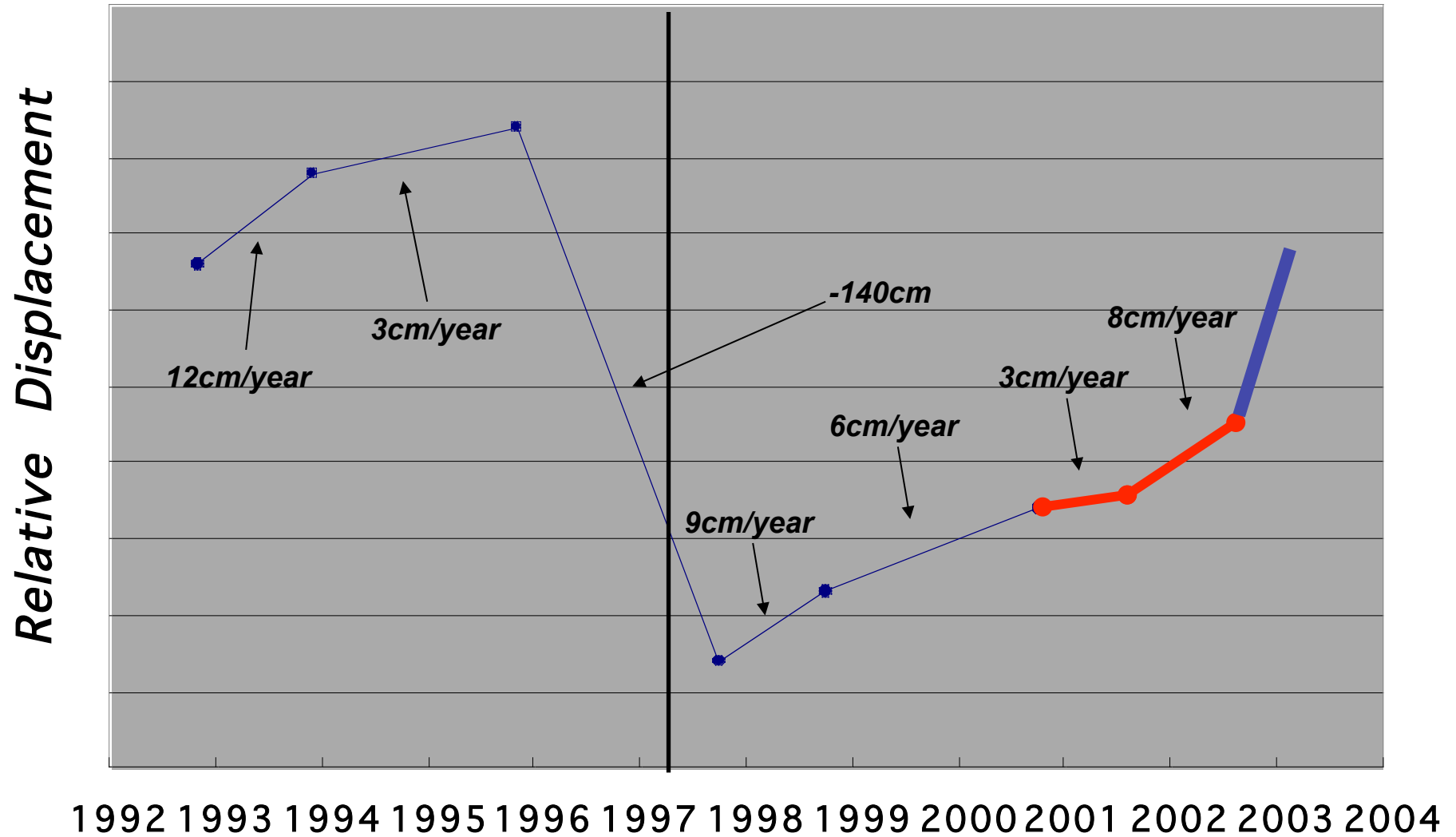
Mann et al.,  
2002

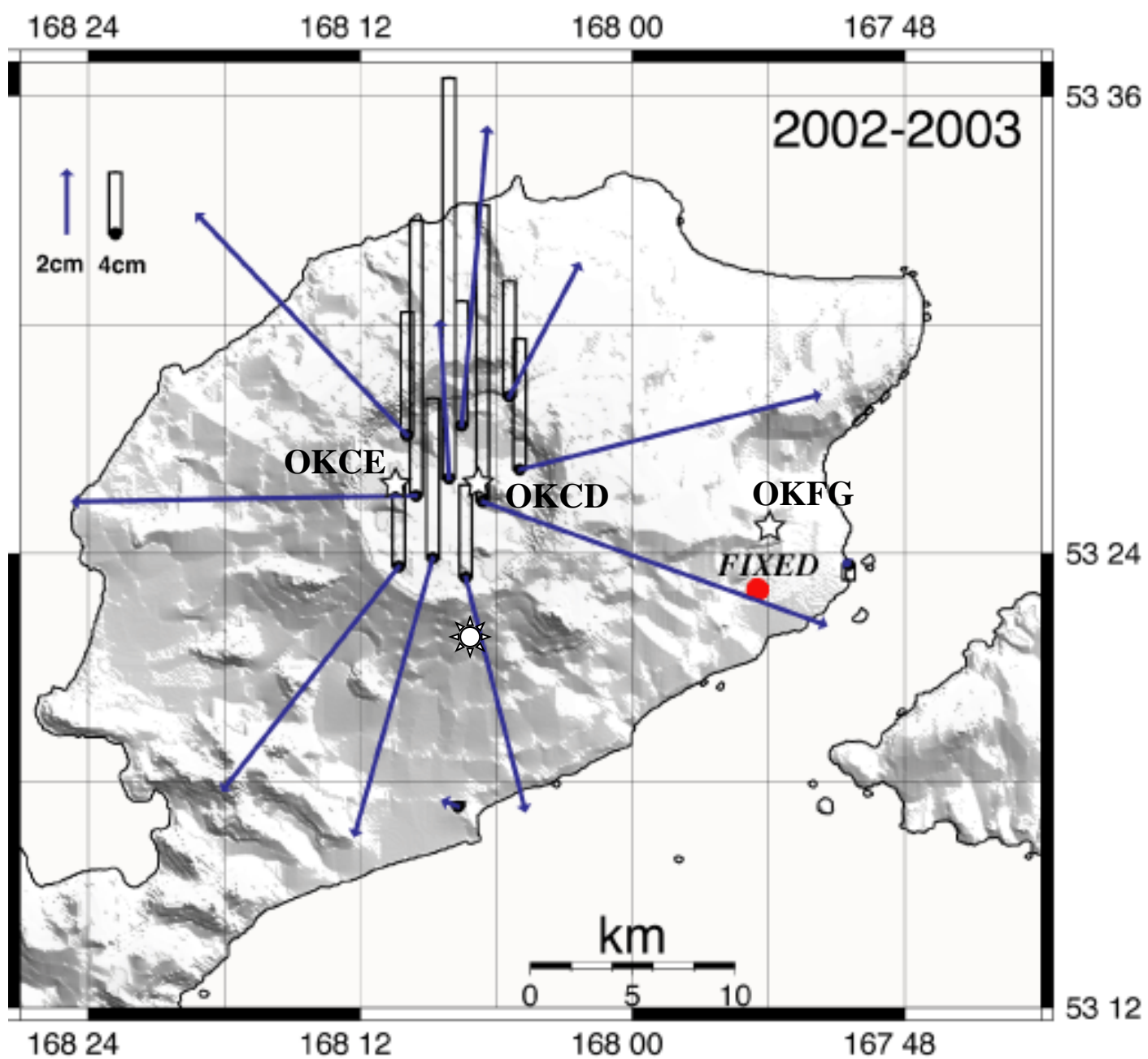
# Source locations and depths



Mann et al., 2002

# Uplift, Combined InSAR+GPS

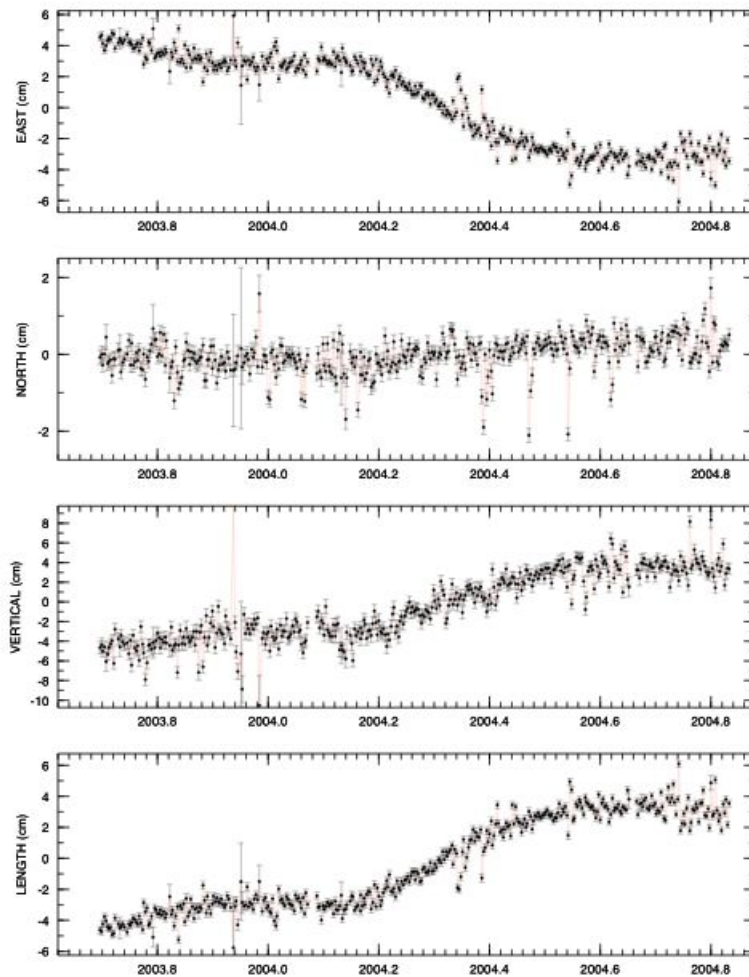




# Intra-caldera Time Series

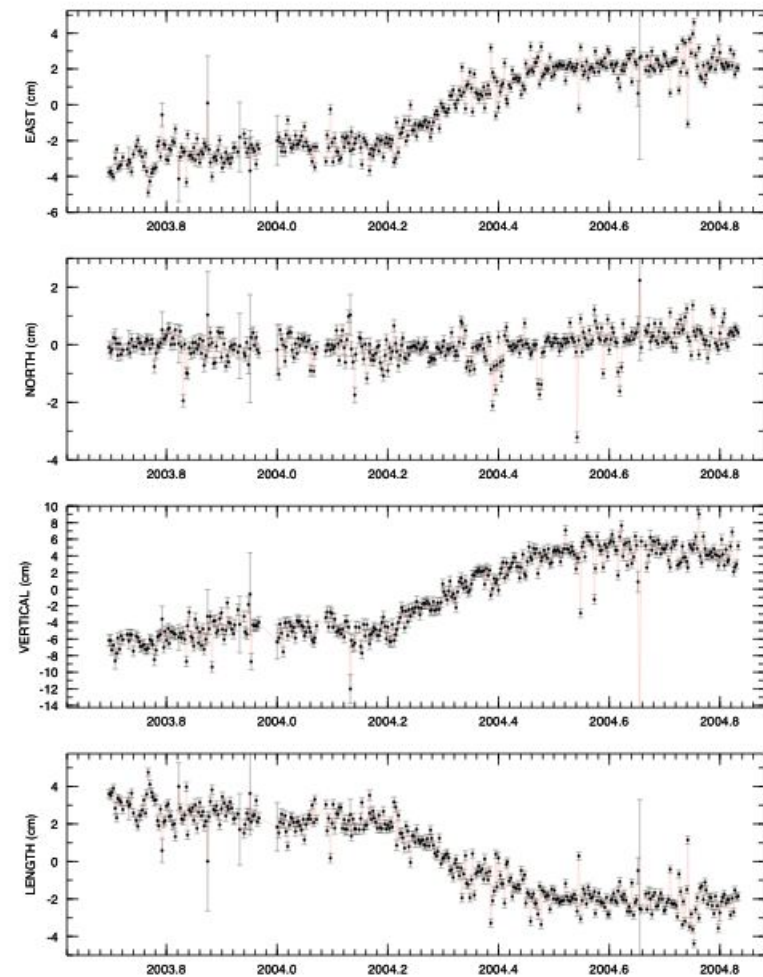
**OKCE relative to OKFG**

17.02 17.02 17.02 km



**OKCD relative to OKFG**

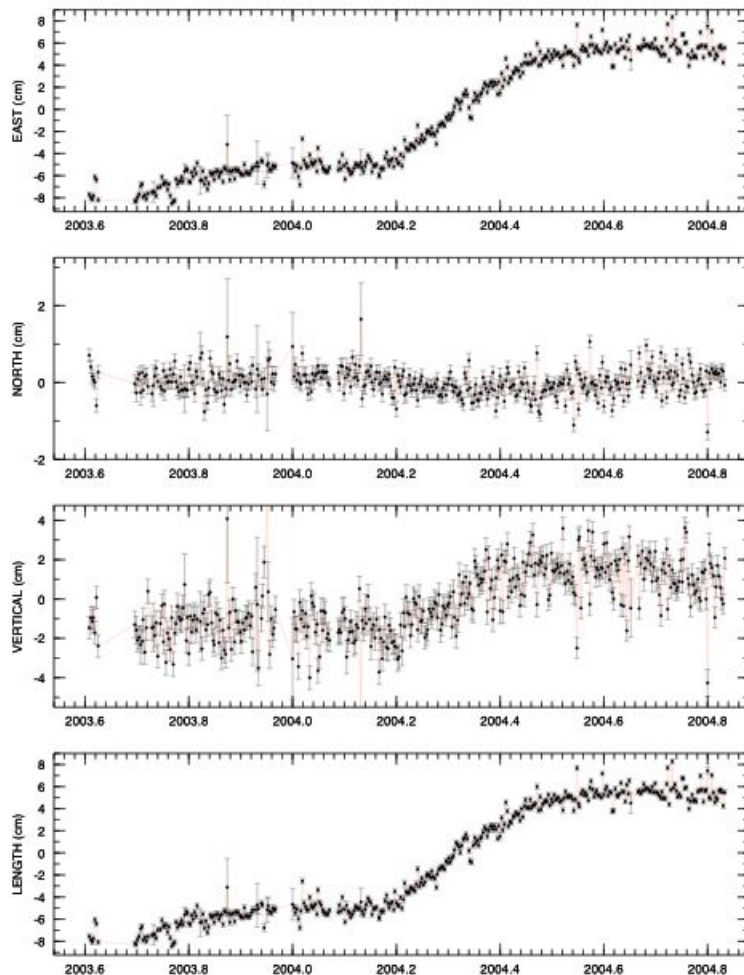
13.65 13.65 13.65 km





# Rapid Inflation Event

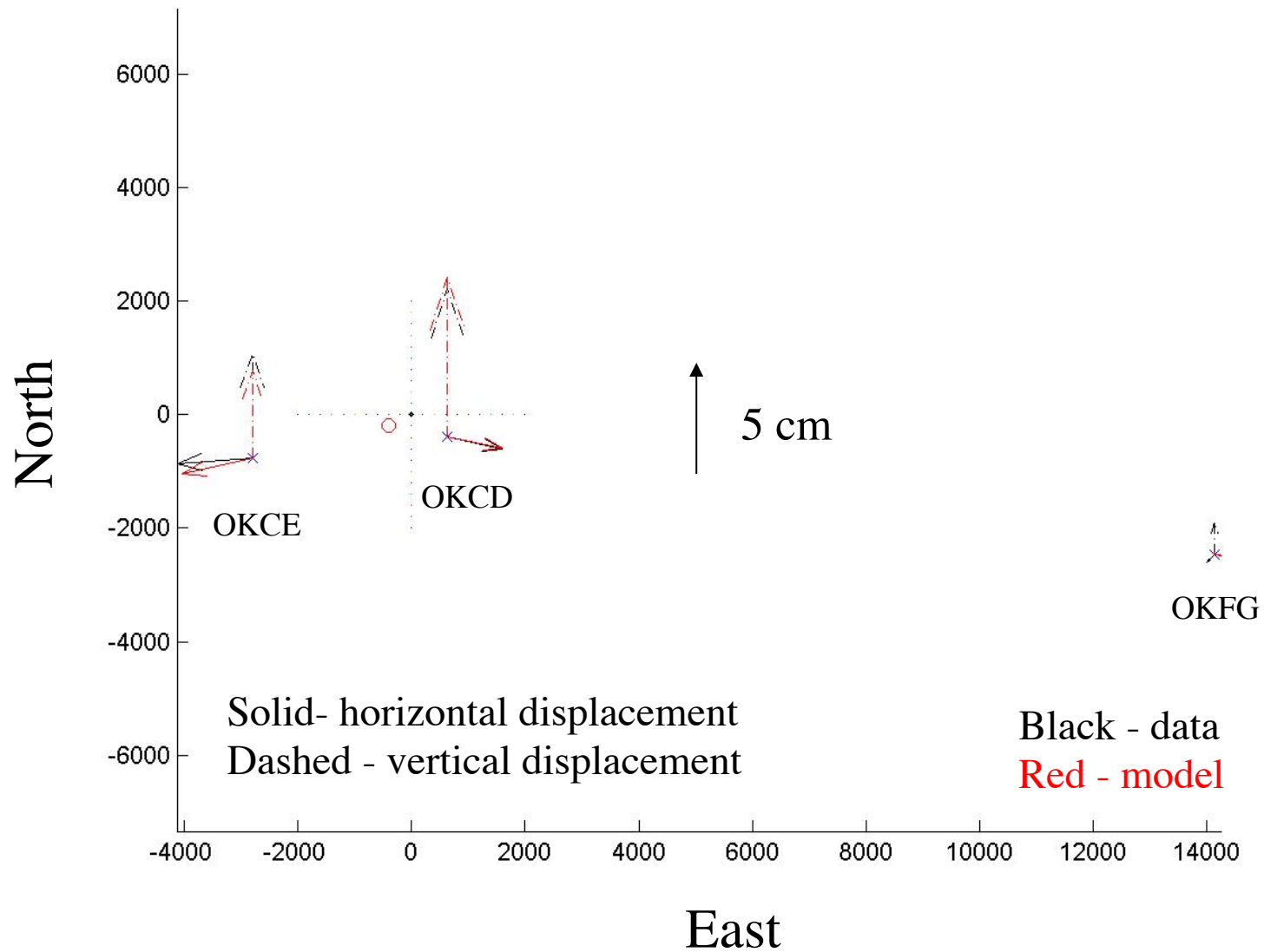
OKCD relative to OKCE  
3.45 3.45 3.45 km



- Inflation occurs almost continuously
- There was a very rapid inflation event Feb-July 2004
- Is source stronger during event? (more magma) Or simply shallower?

# Before Event

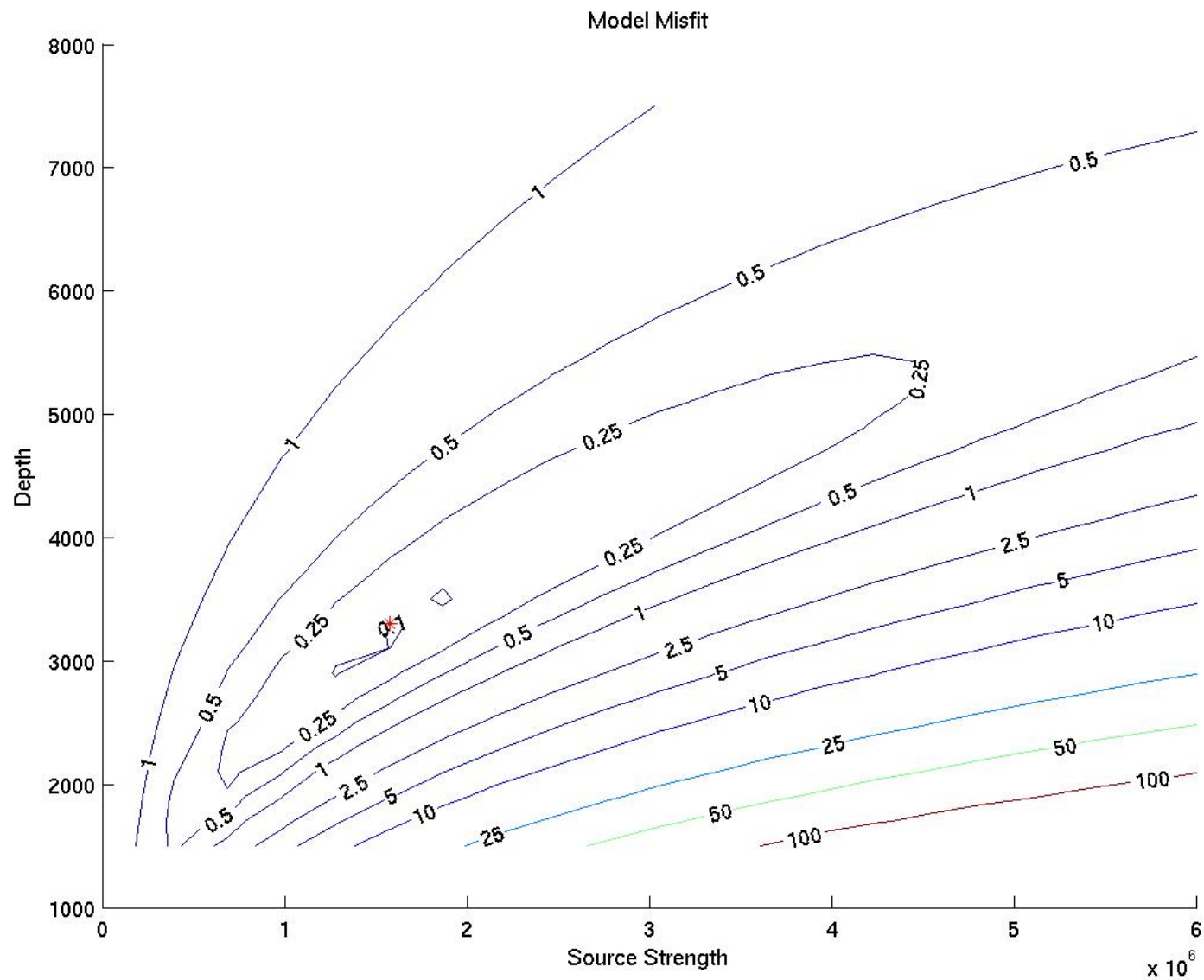
2003.8 to 2004.1



# How Well Determined are Source Parameters?

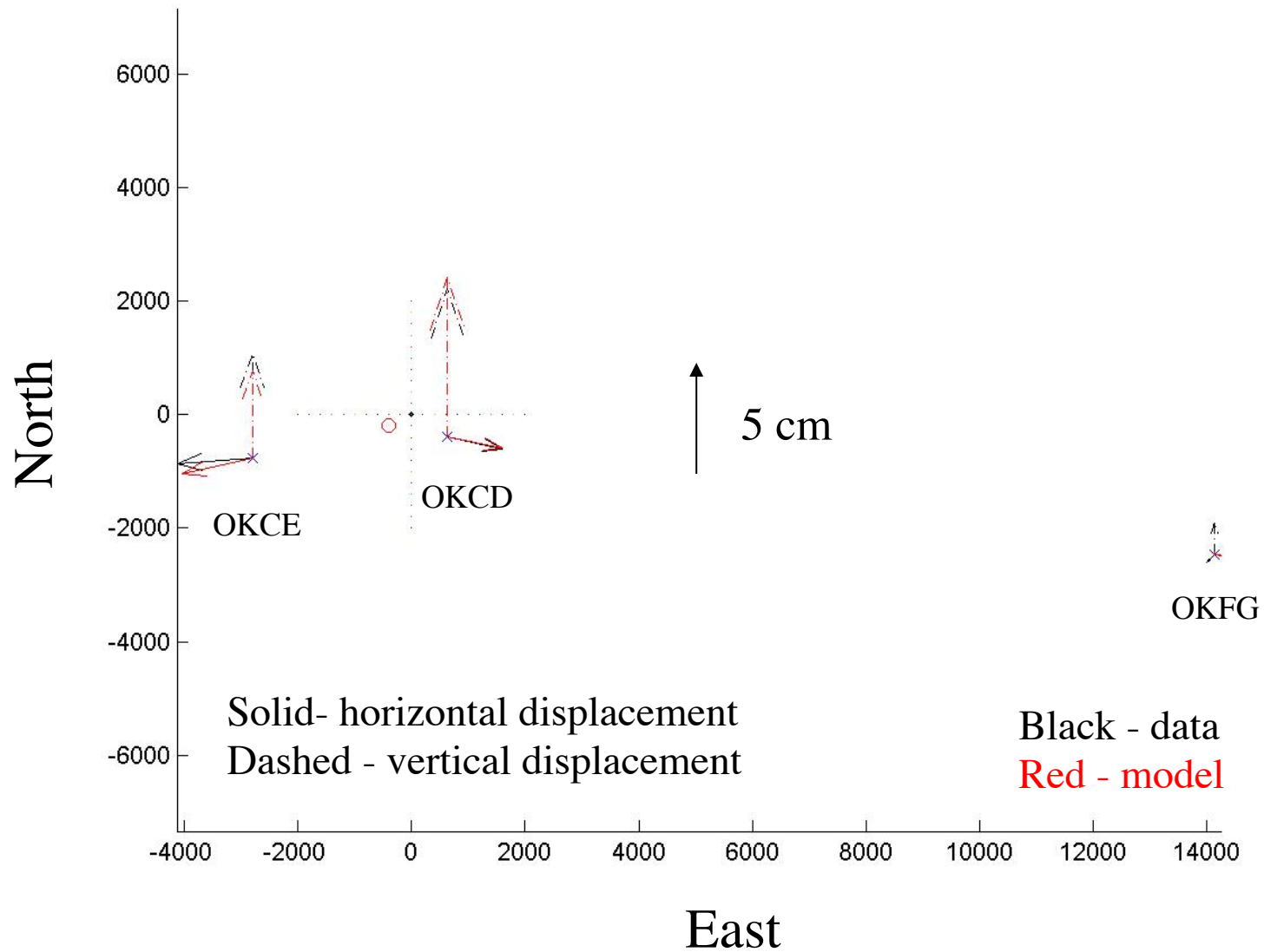
- There is always a tradeoff between depth and source strength ( $p$ , or  $\Delta V$ ). Why?
  - Limitations in sampling
    - Stronger deep source produces same maximum uplift as a weaker shallow source
    - You distinguish the two by having enough data from different distances to distinguish spatial pattern
    - Often don't have enough data, or there are other complications
  - Problem with GPS: only a few displacements
  - Problem in InSAR: often no data directly above source

# Depth-Source Strength Tradeoff



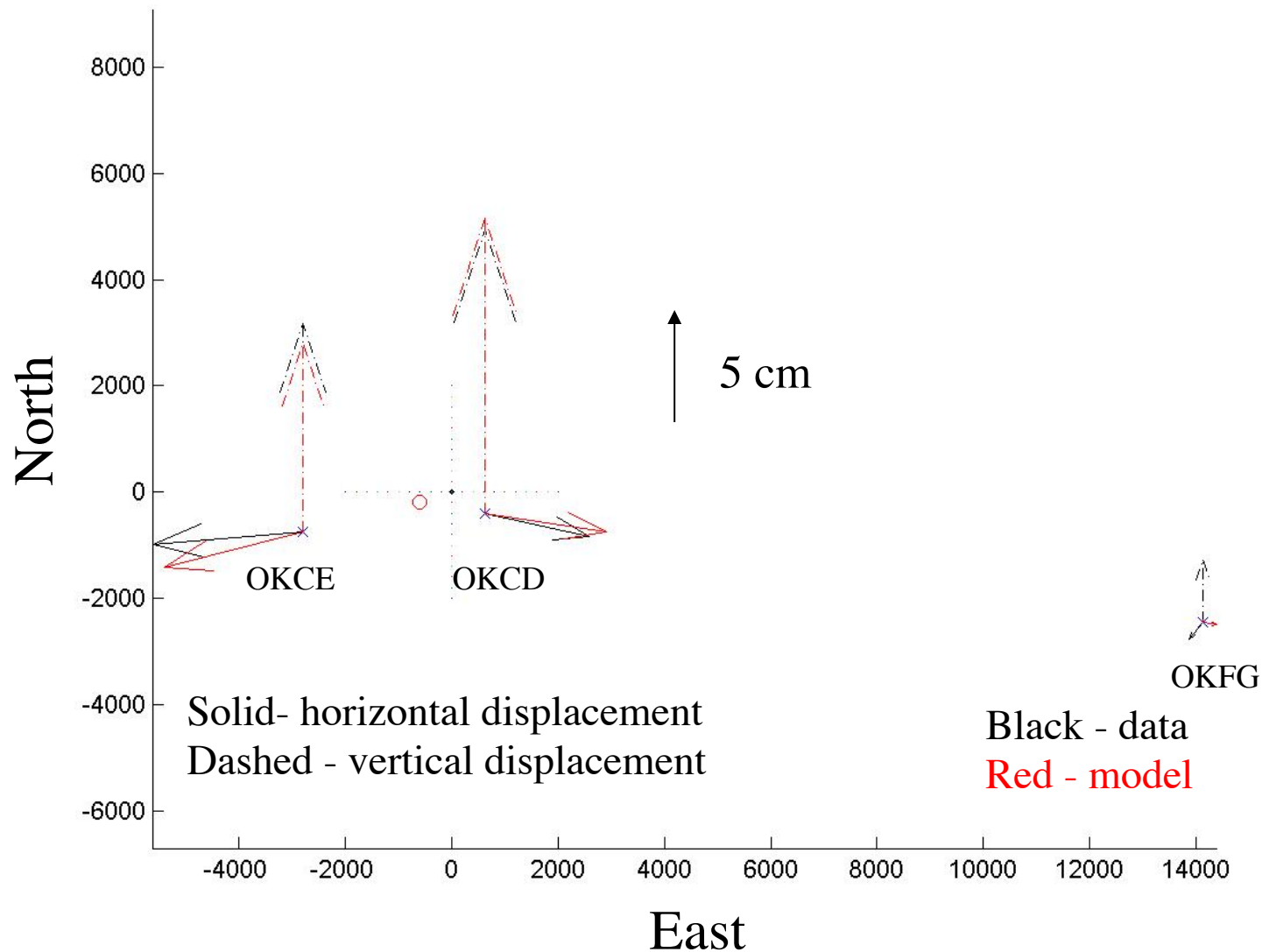
# Before Event

2003.8 to 2004.1



# Displacements During Event

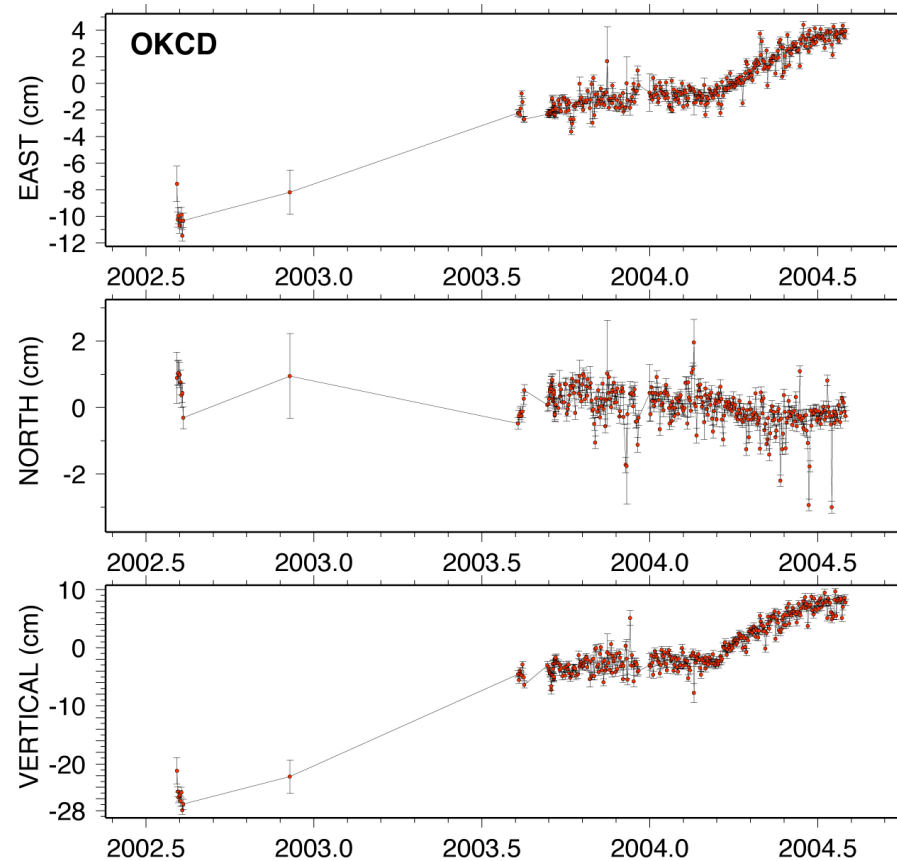
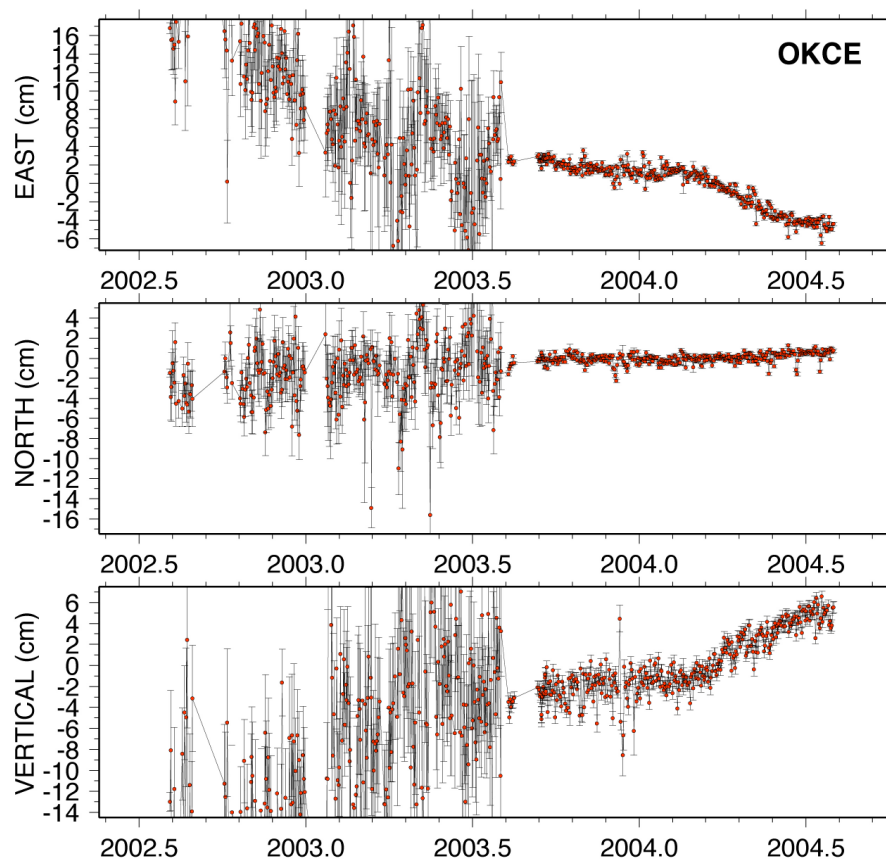
2004.1 to 2004.5



# Comparison of Source Models

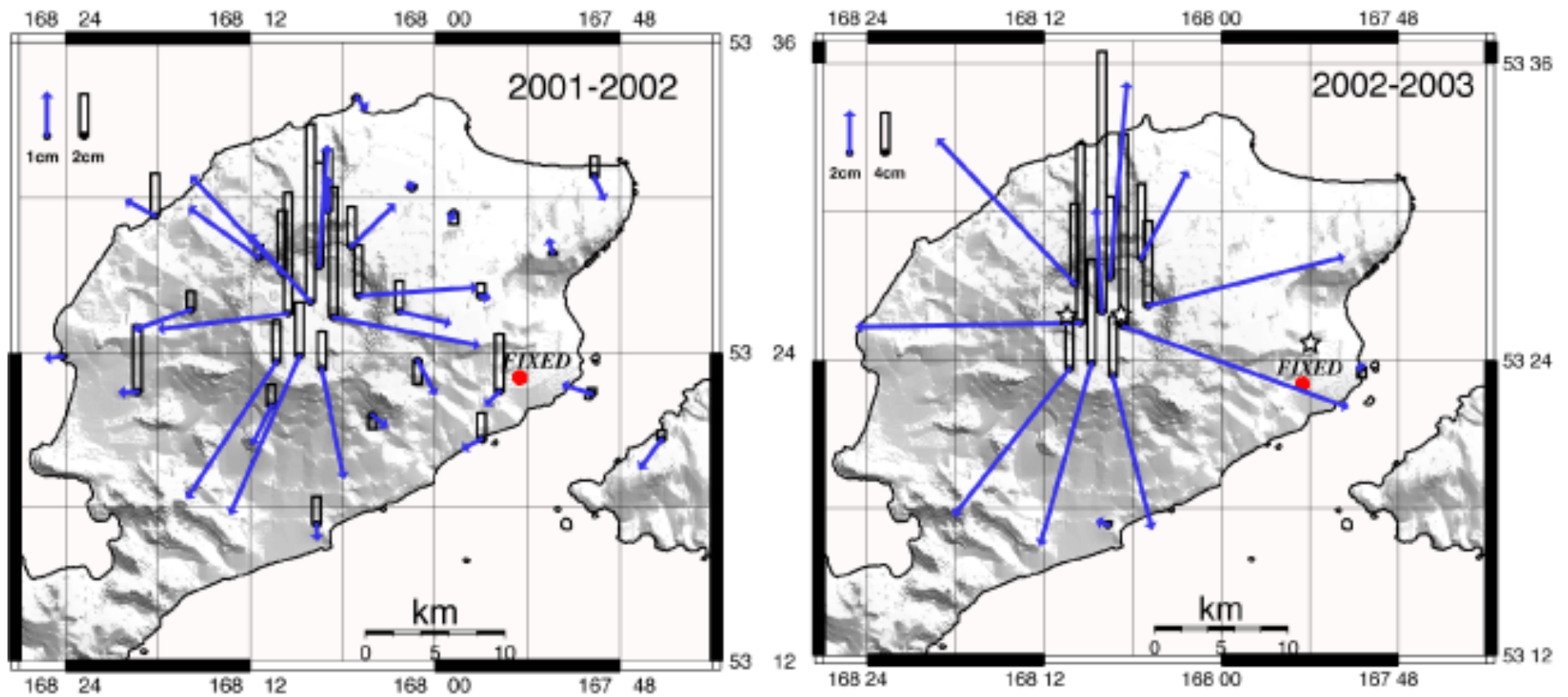
- Horizontal location of source is the same within 200 meters
- Best-fit depth is 3 km for both
- Source strength is 2.1 times higher during event.
  - If we require source strength to be the same during event, it would require source to be at ~2 km, and misfit would be ~3x higher.
  - Volume change *rate* before:  $+7.3 \times 10^6 \text{ m}^3/\text{yr}$  (0.4 yr)
  - Volume change *rate* during:  $+19.6 \times 10^6 \text{ m}^3/\text{yr}$  (0.4 yr)
  - 1990s rate (InSAR):  $\sim +5 \times 10^6 \text{ m}^3/\text{yr}$ ; eruption:  $70 \text{ m}^3$

# Time Series (older data)

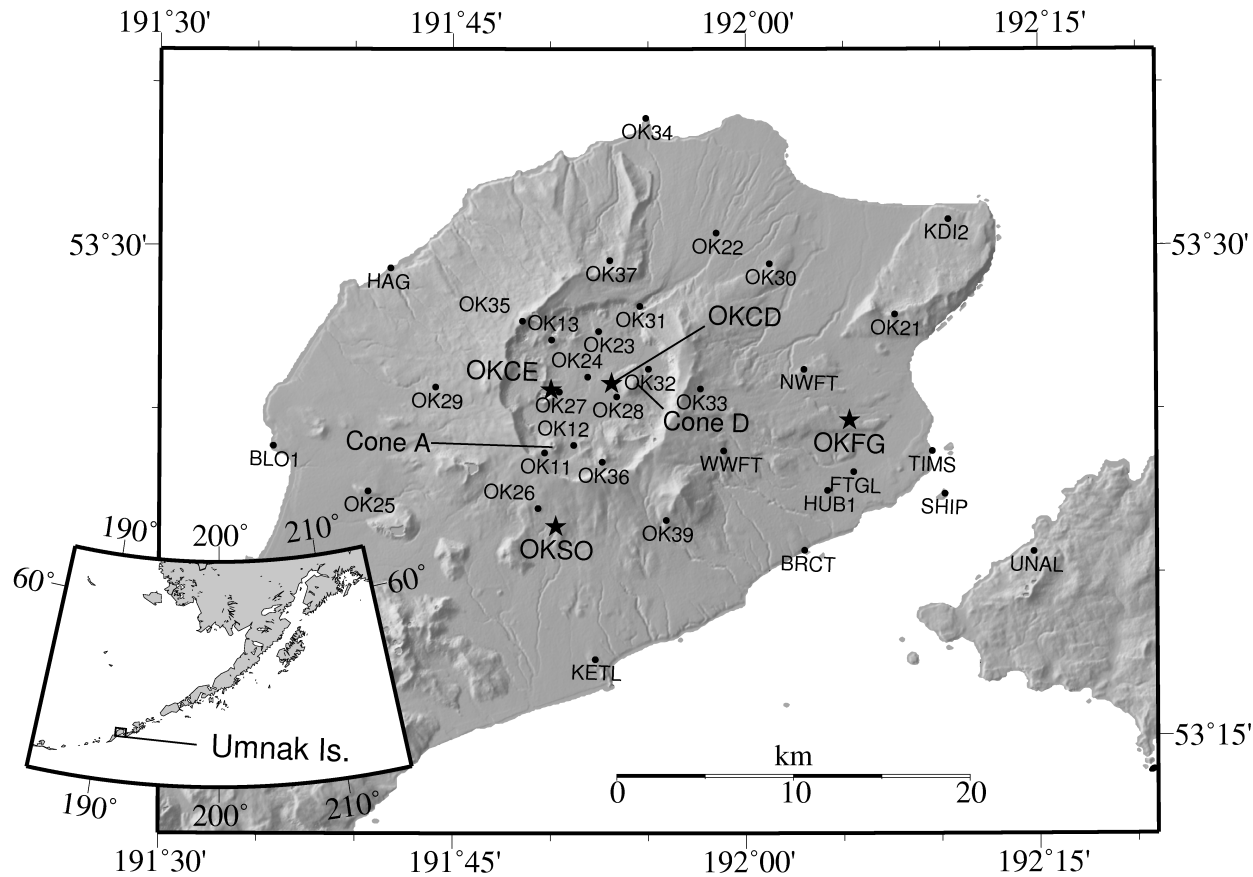




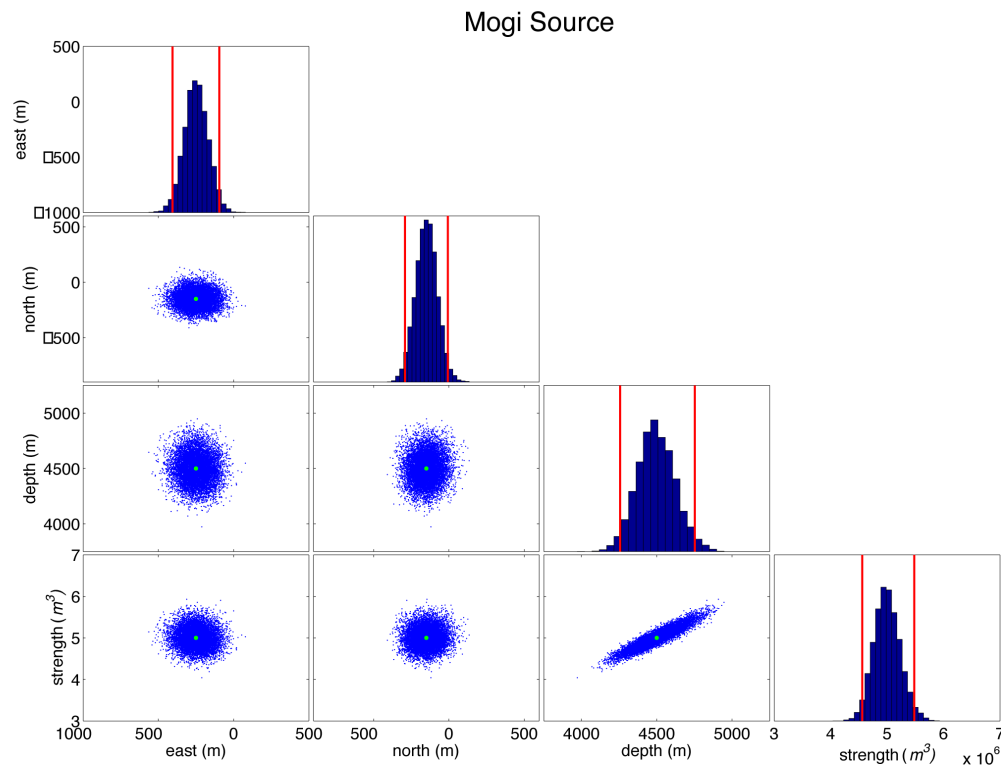
# GPS Displacements



# Okmok Recent Data + Eruption



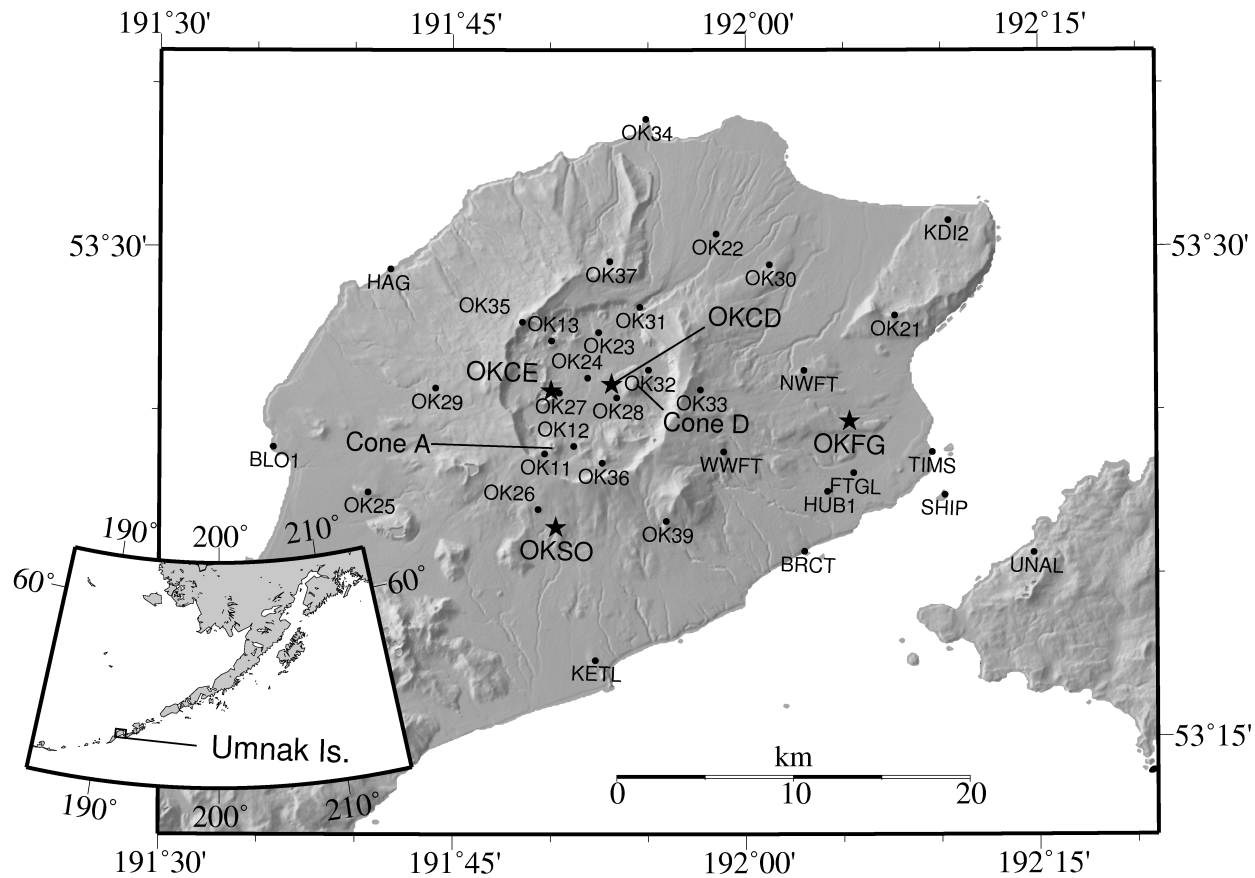
# Theoretical Resolution of Network



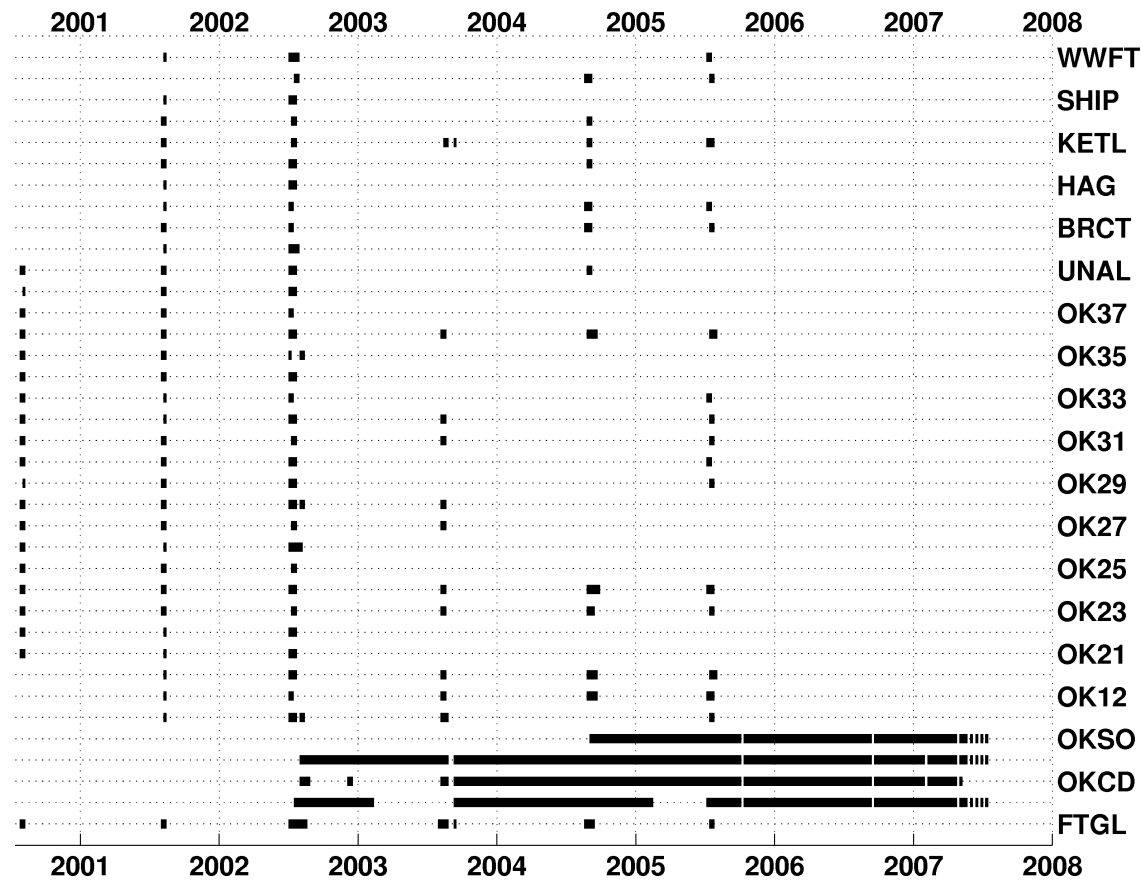
*Fournier et al. [2009, JGR]*

- Test ability of network to determine source location using synthetic data
- Depth-source strength tradeoff is still there even with dense GPS network
  - Can't get rid of it, only minimize it and live with it.

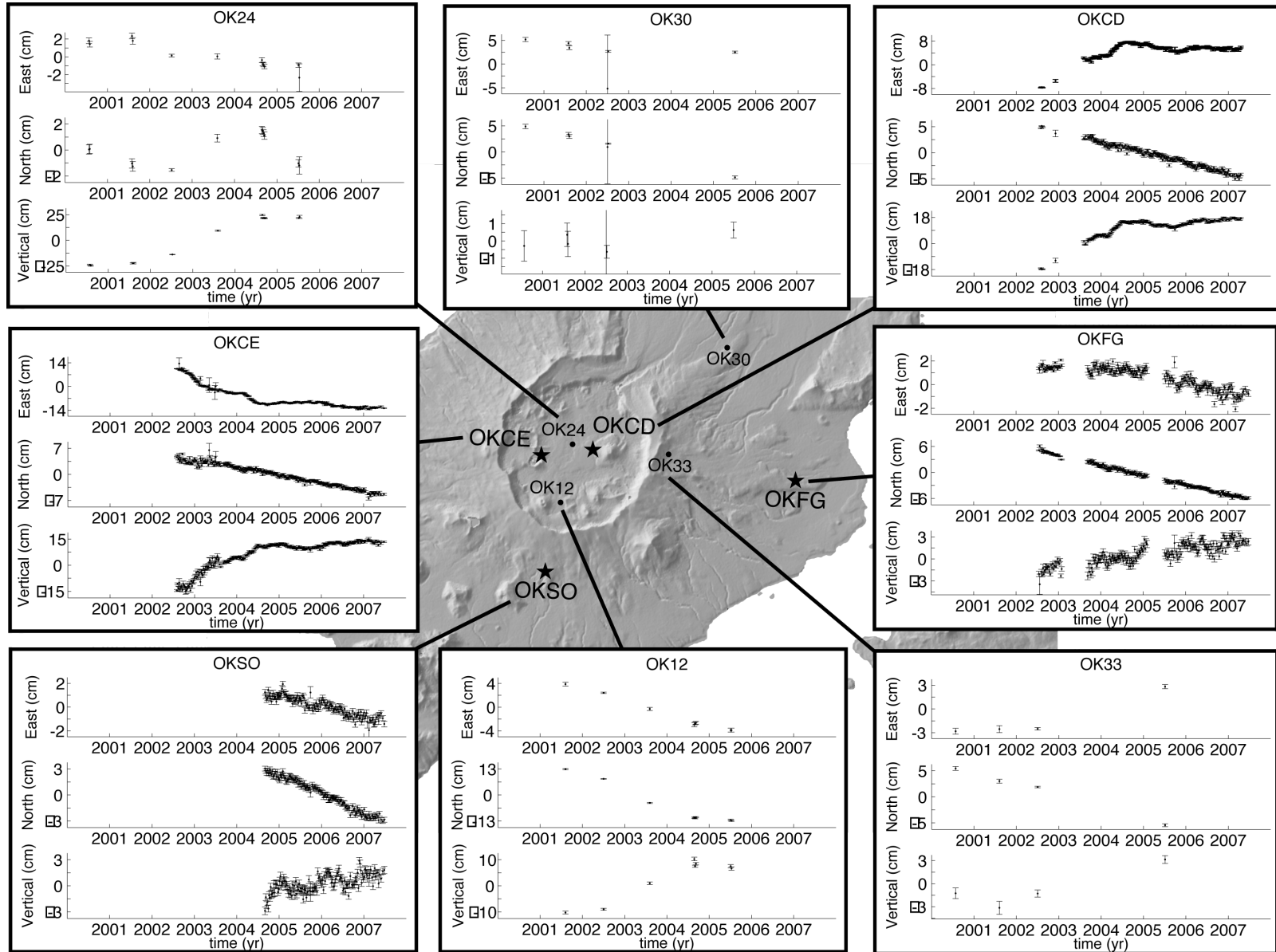
# Okmok GPS Network



# GPS Site Measurement History



# Selected GPS Time Series

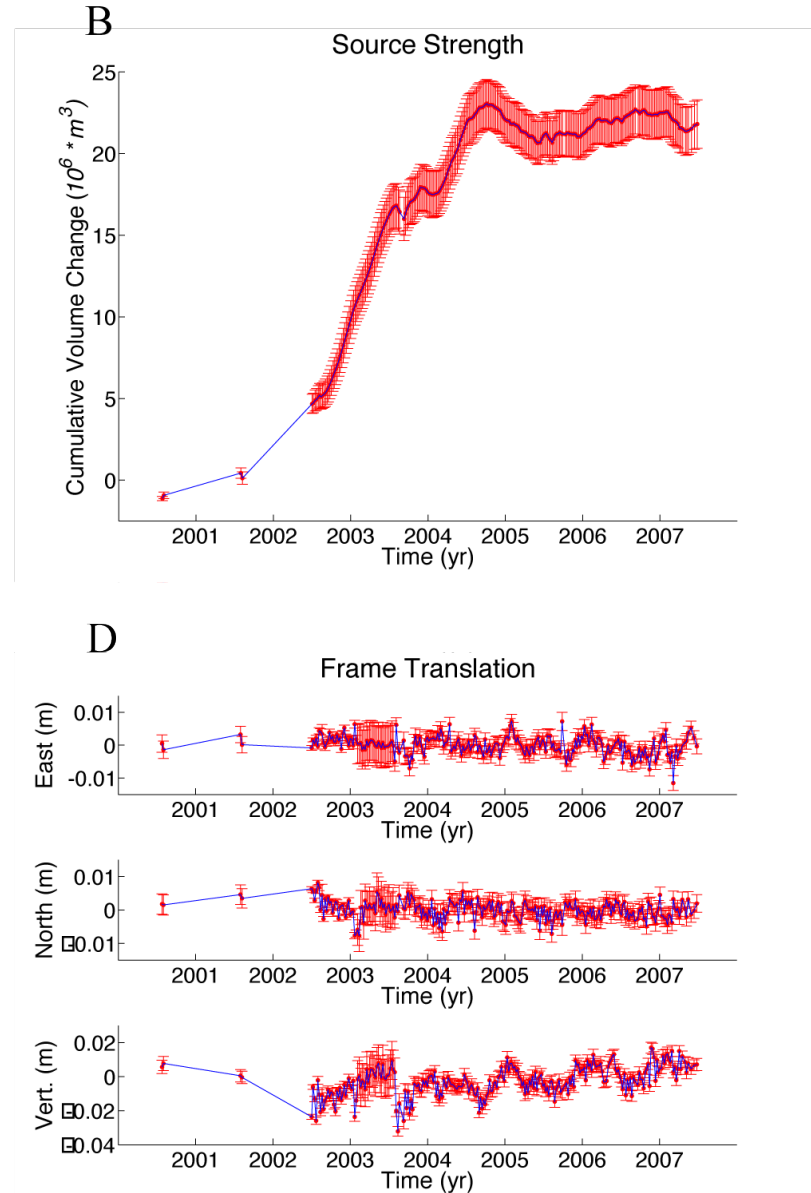
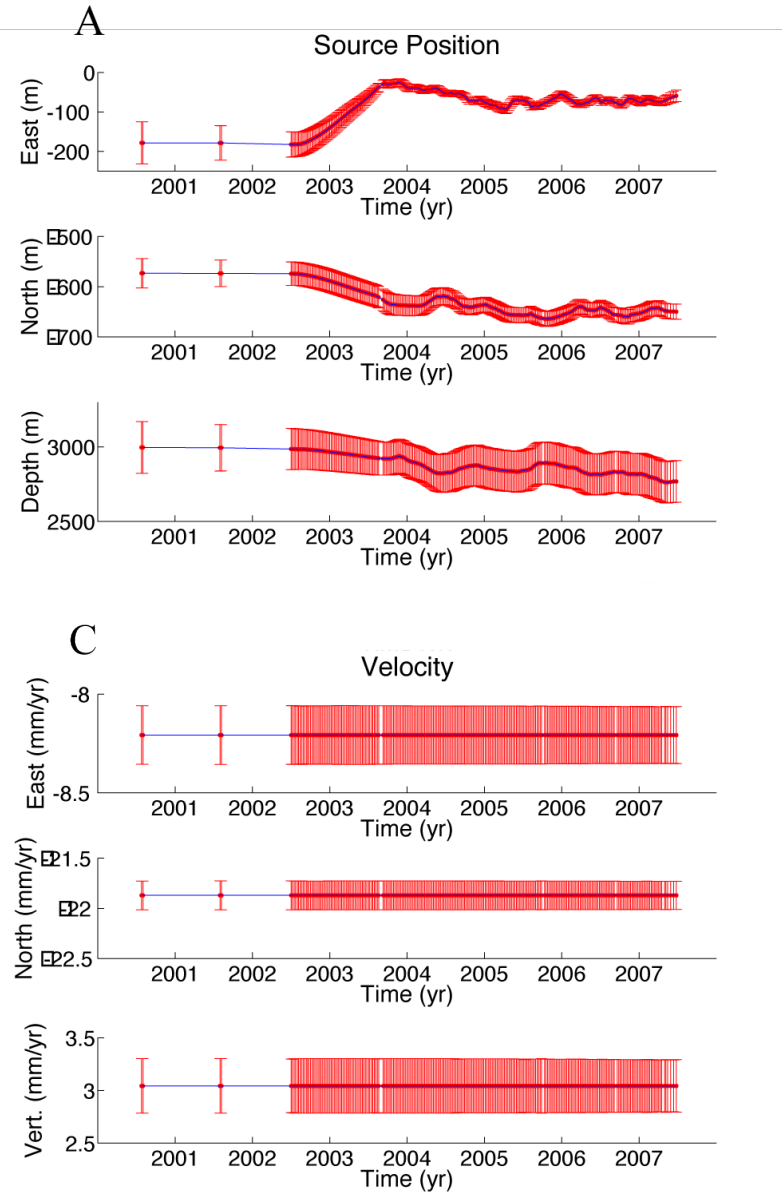




# Modeling Strategy

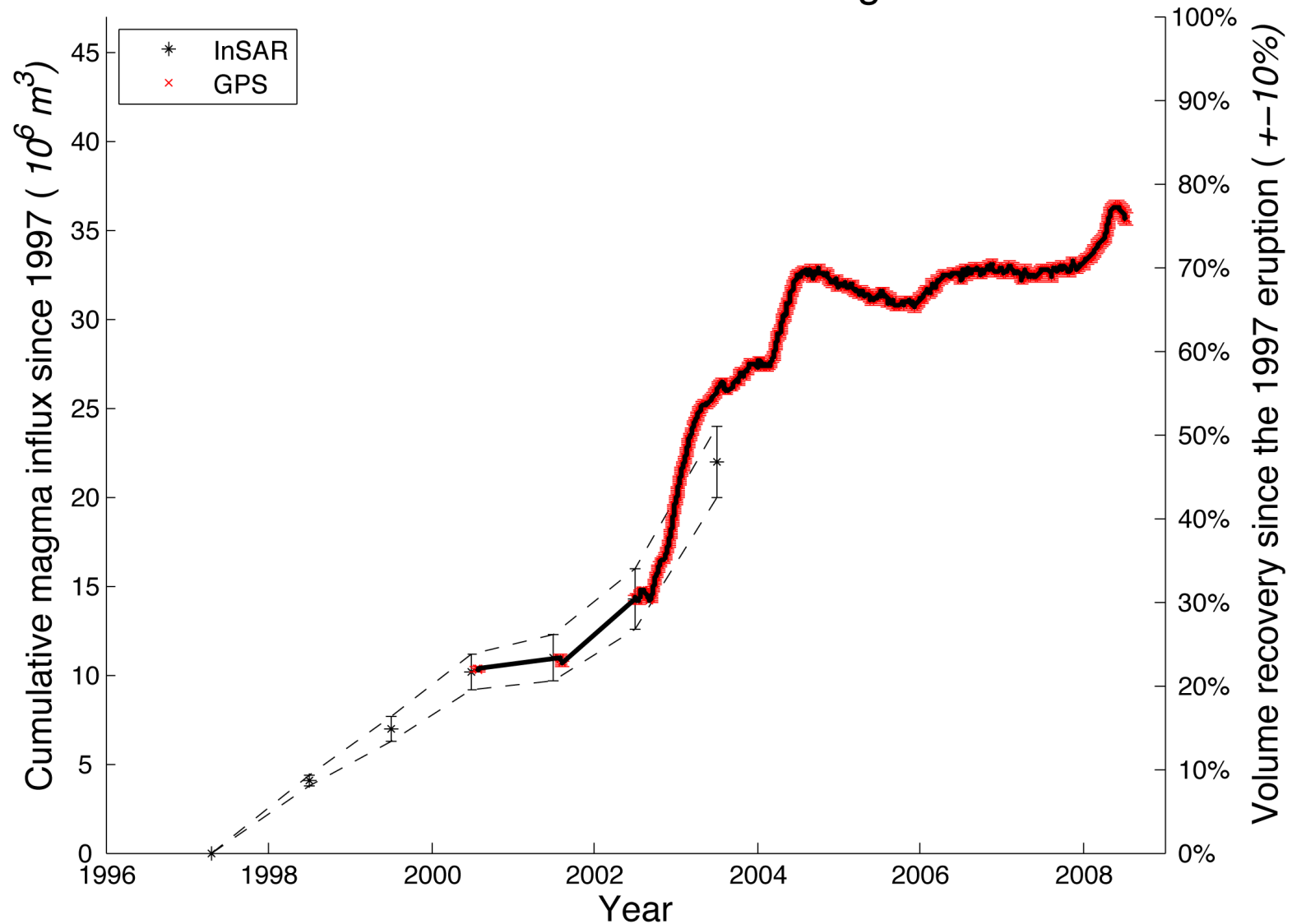
- Use a non-linear Kalman filter approach
  - Kalman filter is a sequential estimation scheme
  - Allows estimates to vary with time according to different stochastic noise models
  - This non-linear version is the “Unscented Kalman Filter”, which attempts to estimate the probability distribution of the parameter values as a function of time.
- Model is Mogi source that is allowed to move with time, plus a steady tectonic velocity for all sites, plus daily reference frame noise.

# Model Results



# Volume Change Inferred from Model

## Cumulative Volume Change

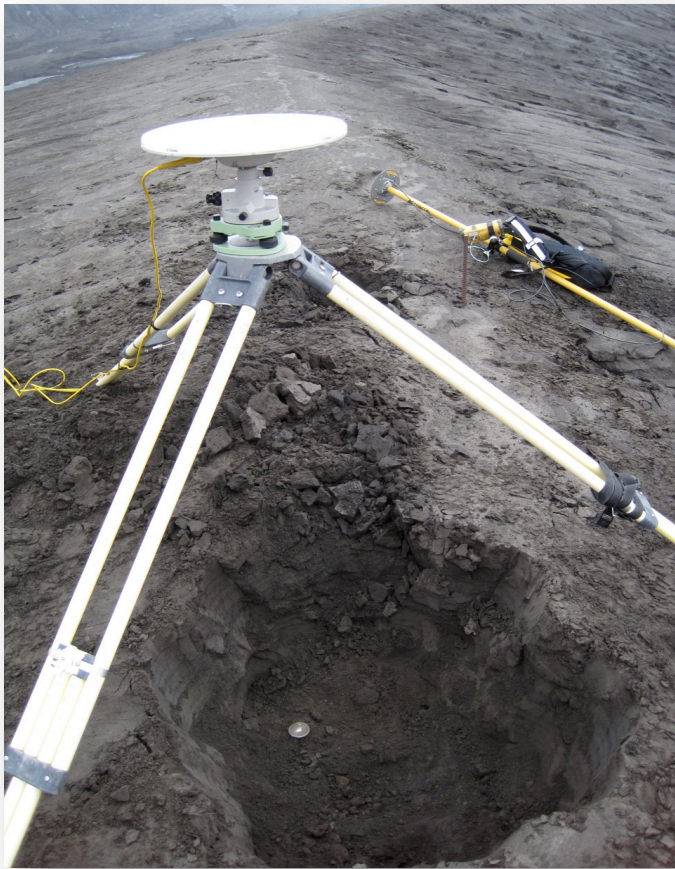


# 2008 Eruption



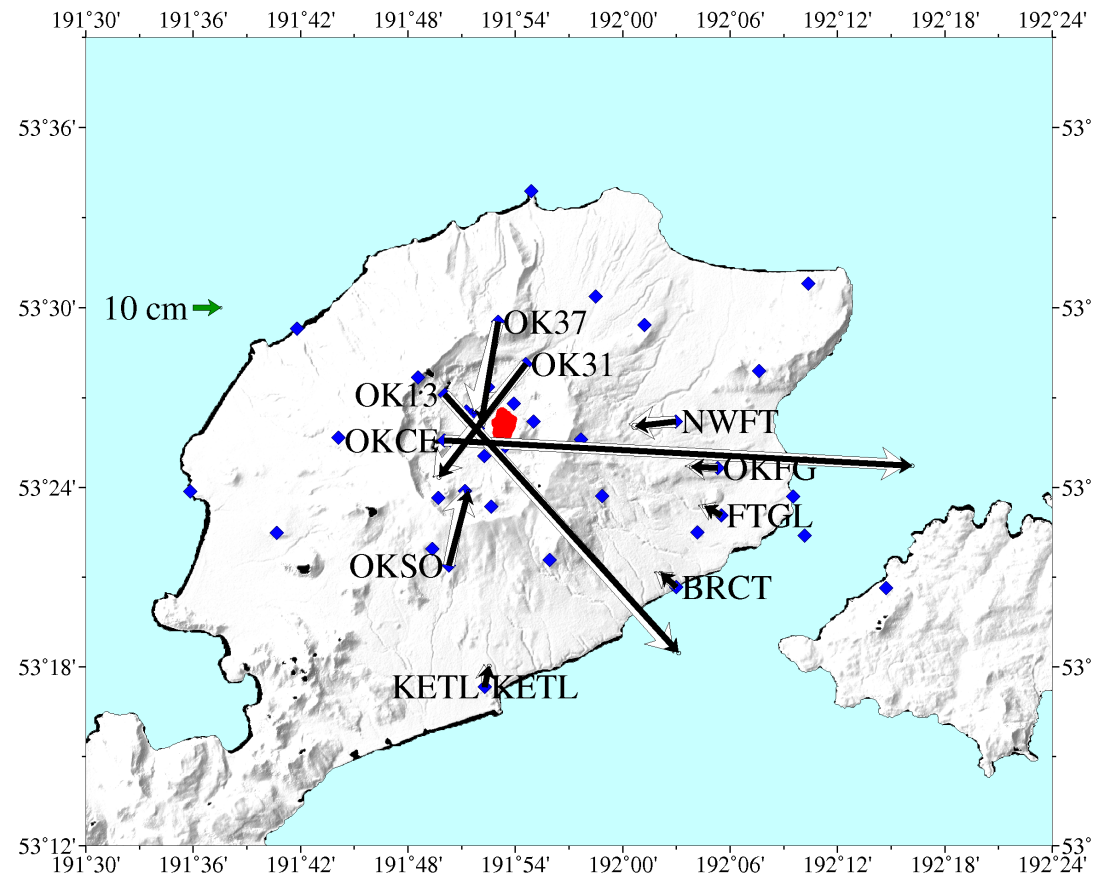


# Inside the Caldera



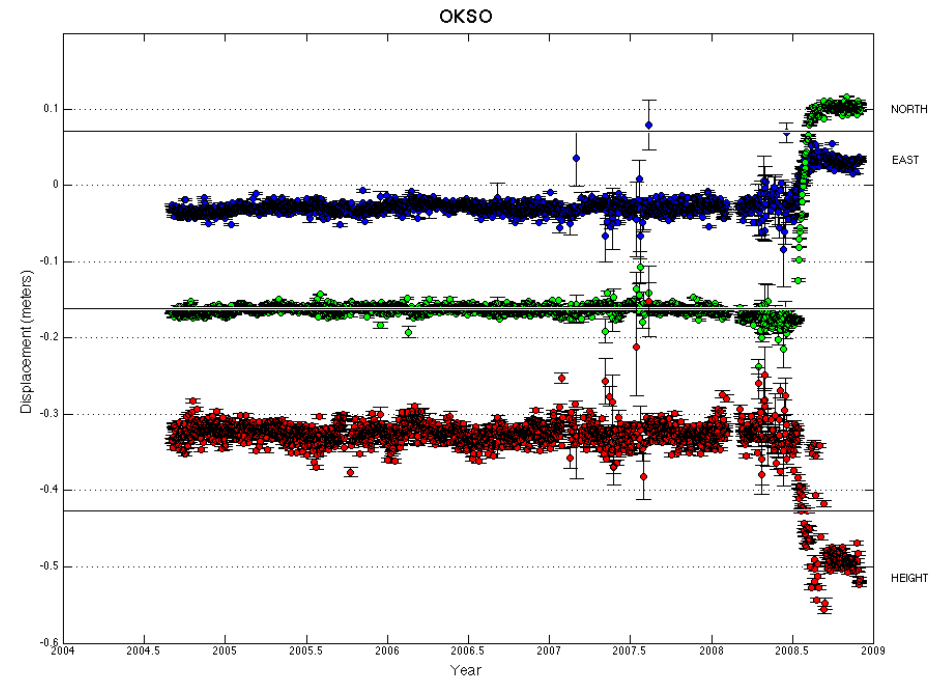
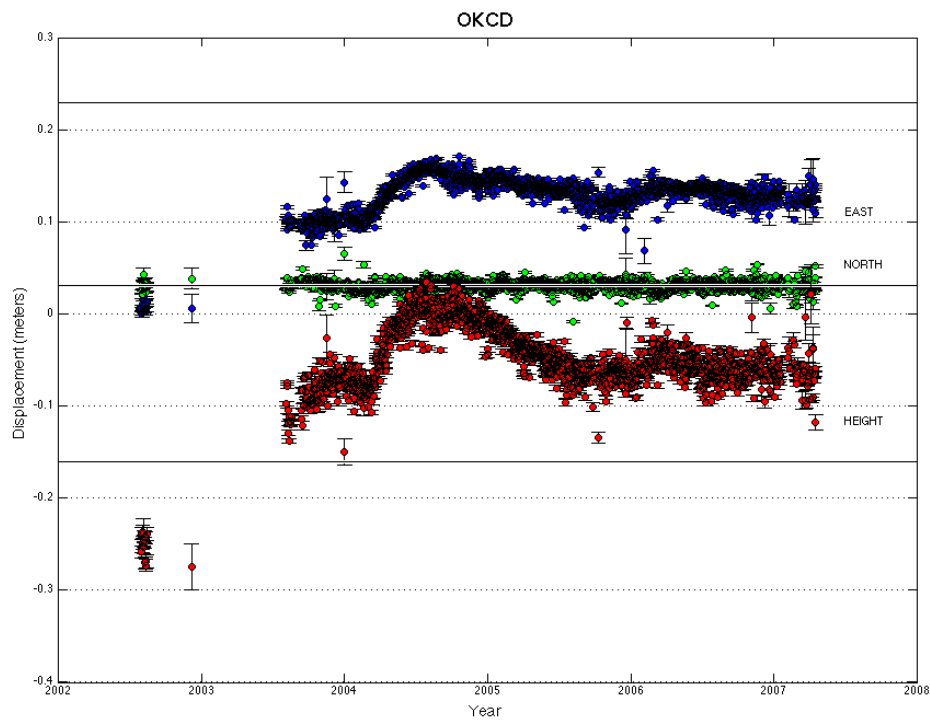
Site OK12 near base of Cone F: 40 cm airfall

# Displacements >1 meter inside caldera





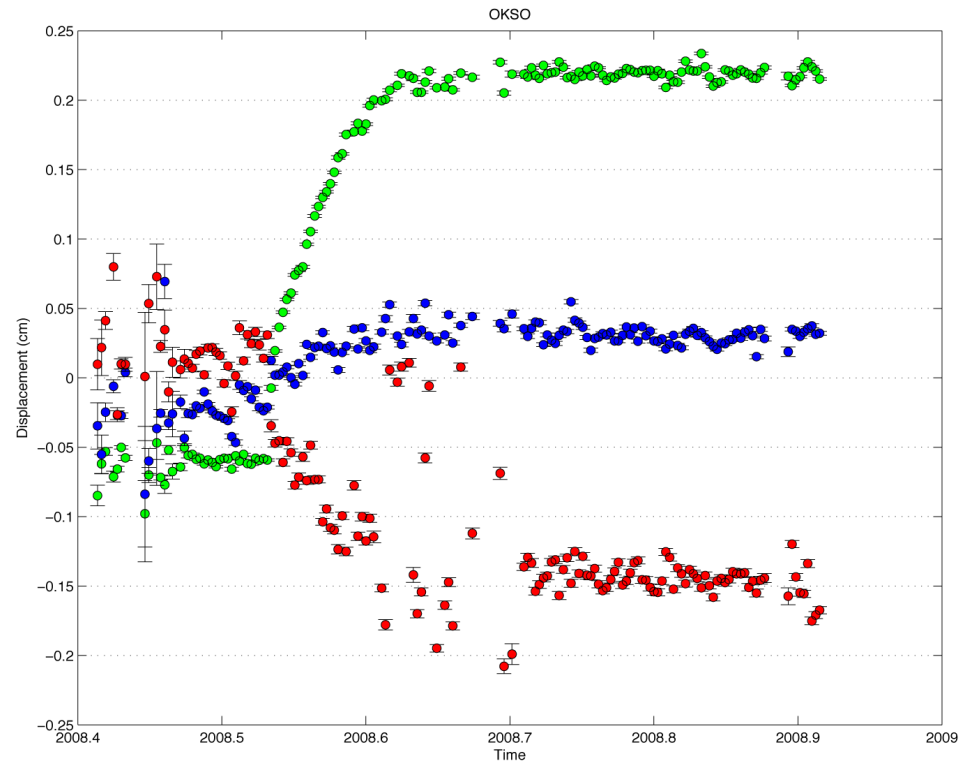
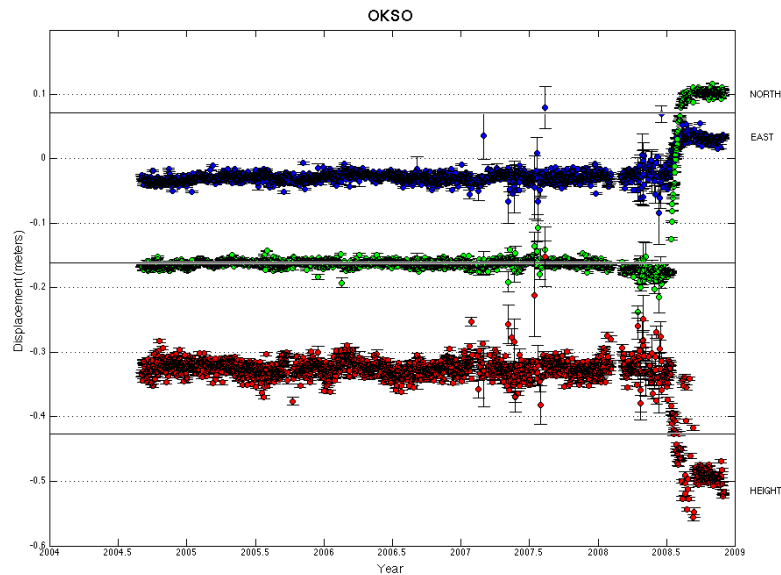
# GPS Site Time Series



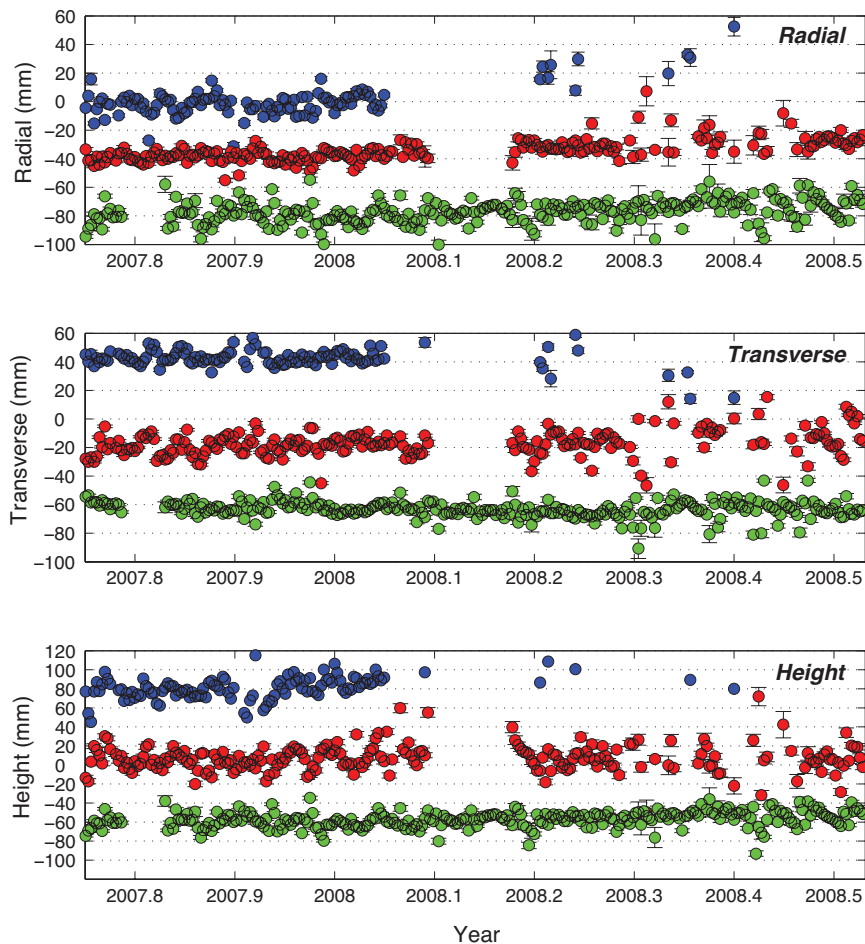
Average rate from 2005.0-2008.0 subtracted

# Eruption Time Period in Detail

*Large scatter in vertical from signals delayed by heavy ash plume.*

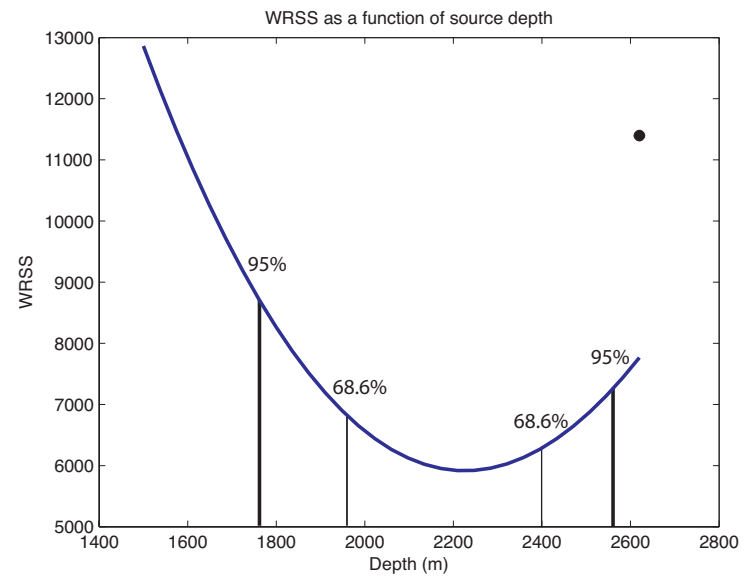
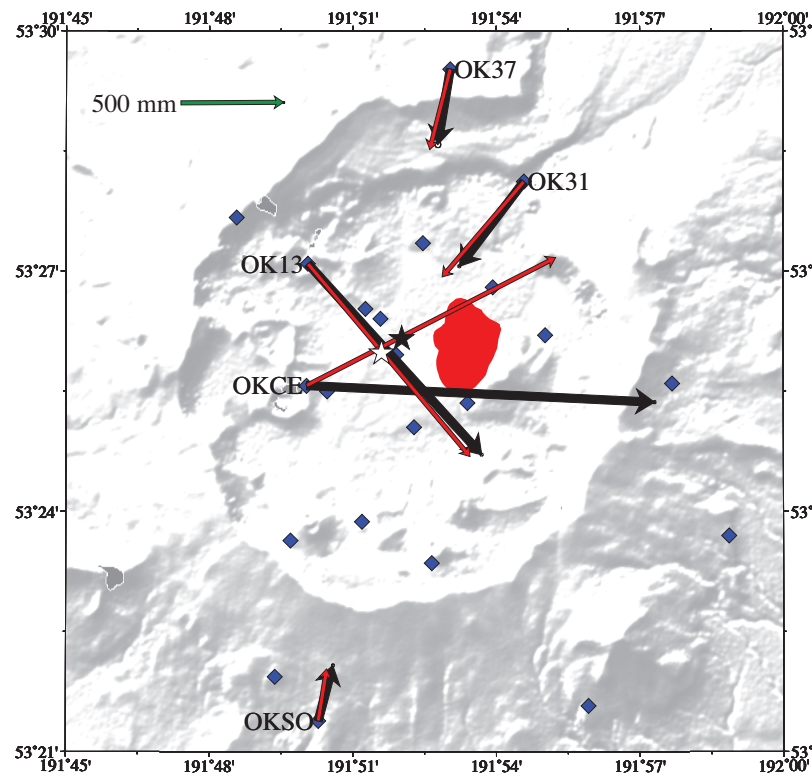


# Final Pre-eruptive Inflation



- Radial, transverse and vertical components relative to the 25-2008 trend.
- Continuation of trend would give a flat line
- OKCE (blue) deviates by early 2008

# Model, Residuals, Misfit



# Deflation History During Eruption

